



Twenty million years of continuous deformation along the Karakorum fault, western Tibet: A thermochronological analysis

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Received 16 October 2005; revised 12 January 2007; accepted 23 March 2007; published 11 July 2007.

[1] The role of the Karakorum fault zone (KFZ) is debated. South of 33°N, ongoing dextral-oblique slip along the SW edge of the Gar basin exhumes metamorphic and magmatic rocks of the Ayilari range. Minerals have recorded a continuum of deformation from temperatures >600–400°C down to <250°C. The ⁴⁰Ar/³⁹Ar ages, the oldest being 21.2 ± 1.0 Ma, yield minimum estimates for the initiation of the KFZ. These are in agreement with the U-Th/Pb ages constraining the onset of deformation at ≥25–22 Ma. Thermochronologic results show slow cooling for the period ~21–14 Ma, followed by rapid cooling between ~14 and 4 Ma. These data demonstrate that right-lateral motion was in progress in the early Miocene and that shear continued at least until 4 Ma, pointing to ≥20 Ma of deformation along the fault. Greenschist facies deformation superimposed upon the medium- to high-grade deformation marks a kinematic change from pure dextral to dextral-normal motion associated with the onset of rapid cooling. At the regional scale, the coexistence of transtension in the Gar basin with transpression documented along the Pangong range farther north suggests another example of the “zipper tectonics” model developed along the Red River fault. The kinematic shift induced the rise of the Ayilari range starting at 16–12 Ma and the incision of major river courses. The Indus River might have become captive of the relief at this time. The river’s 120 km of apparent offset implies dextral motion at a long-term

rate of $\geq 8.5 \pm 1.5 \text{ mm yr}^{-1}$. **Citation:** Valli, F., N. Arnaud, P. H. Leloup, E. R. Sobel, G. Mahéo, R. Lacassin, S. Guillot, H. Li, P. Tapponnier, and Z. Xu (2007), Twenty million years of continuous deformation along the Karakorum fault, western Tibet: A thermochronological analysis, *Tectonics*, 26, TC4004, doi:10.1029/2005TC001913.

1. Introduction

[2] The understanding of how strain is partitioned between shortening and lateral extrusion during continent-continent collision hinges on the determination of the geometry, kinematics, slip rate, and life span of major crustal-scale faults. In the Indo-Asian collision, it is clear that thrusting in the Himalaya absorbs a significant portion of the northward indentation of India through north-south shortening. It remains debated, however, whether strike-slip shear zones are long-lasting, lithospheric features, steadily accommodating large amounts of strain during long time spans [e.g., Tapponnier *et al.*, 2001], or short-lived transient features, restricted to the crust, that contribute to distributed regional deformation [e.g., Molnar and Houseman, 2004].

[3] In southern and western Tibet, north of the Himalayas, the 1000 km long dextral Karakorum Fault zone (KFZ) may be a key to assess this issue (Figure 1). Published estimates of the age of initiation of the KFZ vary greatly between ≥11 Ma [Matte *et al.*, 1996], <5 Ma [Searle, 1996], <18 Ma [Searle *et al.*, 1998, and references therein], <13 Ma [Murphy *et al.*, 2000, 2002], ≥23 Ma [Lacassin *et al.*, 2004], and between 13.7 ± 0.3 and 15.7 ± 0.5 Ma [Phillips *et al.*, 2004]. Long-term slip rate estimates averaged on the Pleistocene to present time period also vary over more than one order of magnitude, from ~30 mm yr⁻¹ [Avouac and Tapponnier, 1993; Liu, 1993; Matte *et al.*, 1996], 10.7 ± 0.7 mm yr⁻¹ [Chevalier *et al.*, 2005], to ~4 mm yr⁻¹ [Brown *et al.*, 2002; Murphy *et al.*, 2002]. A slip rate of 10 ± 3 mm yr⁻¹ was averaged for the entire life span of the fault [Lacassin *et al.*, 2004].

[4] The present article focuses on the Ayilari range west of the Gar-Baer pull-apart basin, where metamorphic rocks deformed by the KFZ are locally exhumed. The ⁴⁰Ar/³⁹Ar dating on syntectonic minerals yield estimates for the minimum initiation age of the KFZ and provide indications of continuous fault activity. Together with fission track and (U-Th)/He ages on apatite, they provide the time-temperature

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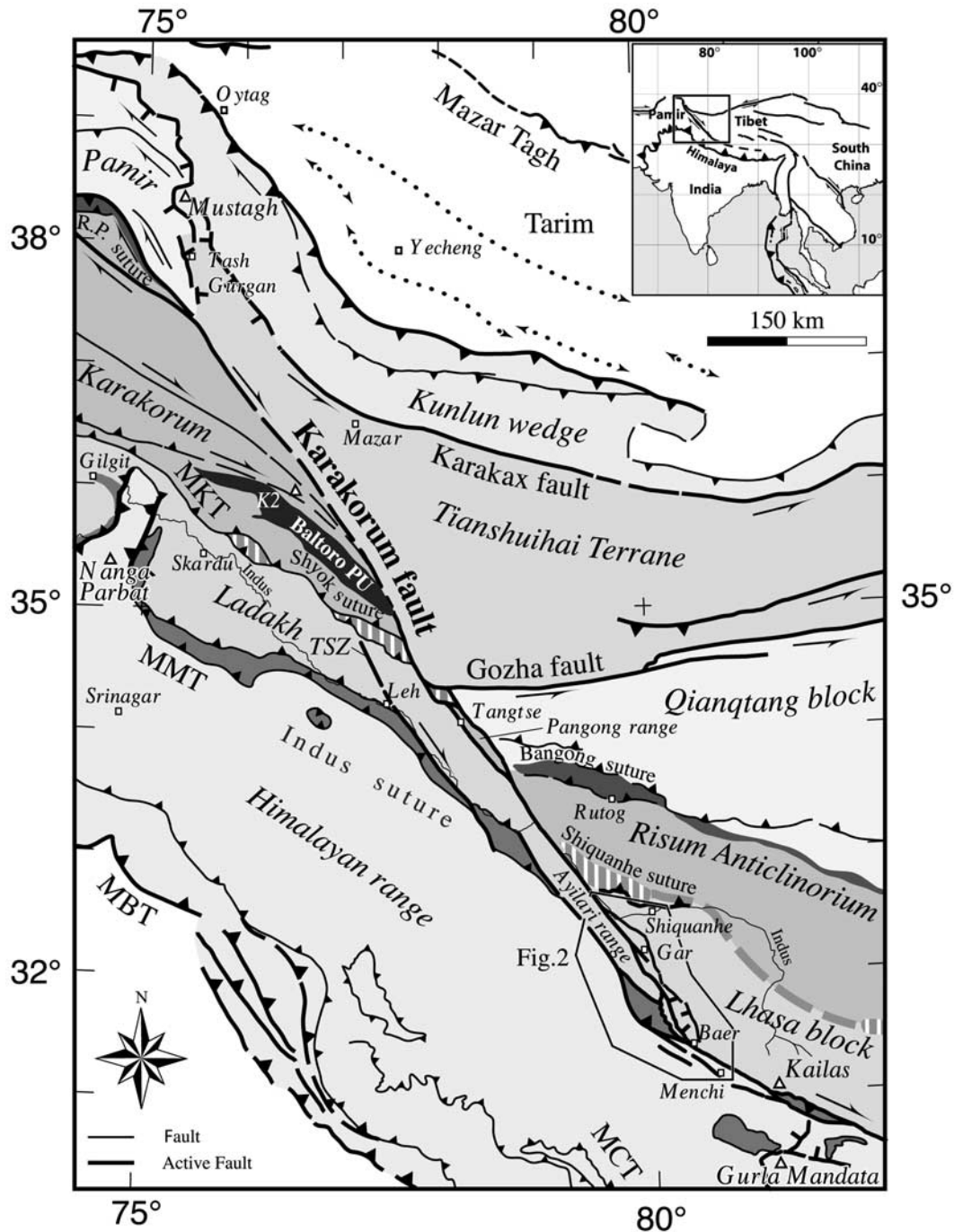


Figure 1. Large-scale sketch map of KFZ and adjacent units. Inset shows location in the large-scale Asian active tectonic framework. Dark gray areas outline the Indus-Zangbo, the Bangong-Nujiang/Rushan-Pshart suture zones, and the main continental blocks. Hatched tones correspond to Shiquanhe and Shyok suture zones. Faults are mapped from Brunel *et al.* [1994], Lacassin *et al.* [2004], Ministry of Geology of USSR [1989], Tapponnier *et al.* [2001], and Weinberg and Dunlap [2000]. PU, Plutonic Unit; TSZ, Thanglasgo Shear Zone. Map is projected in UTM 44 using the ellipsoid WGS84.

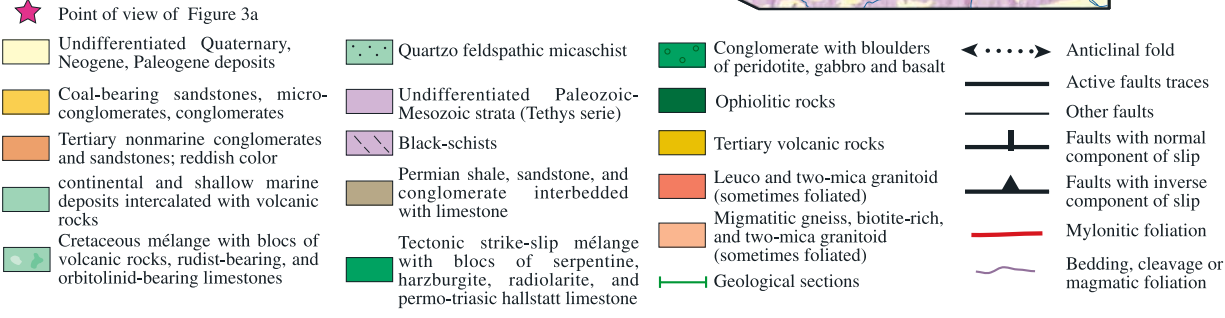
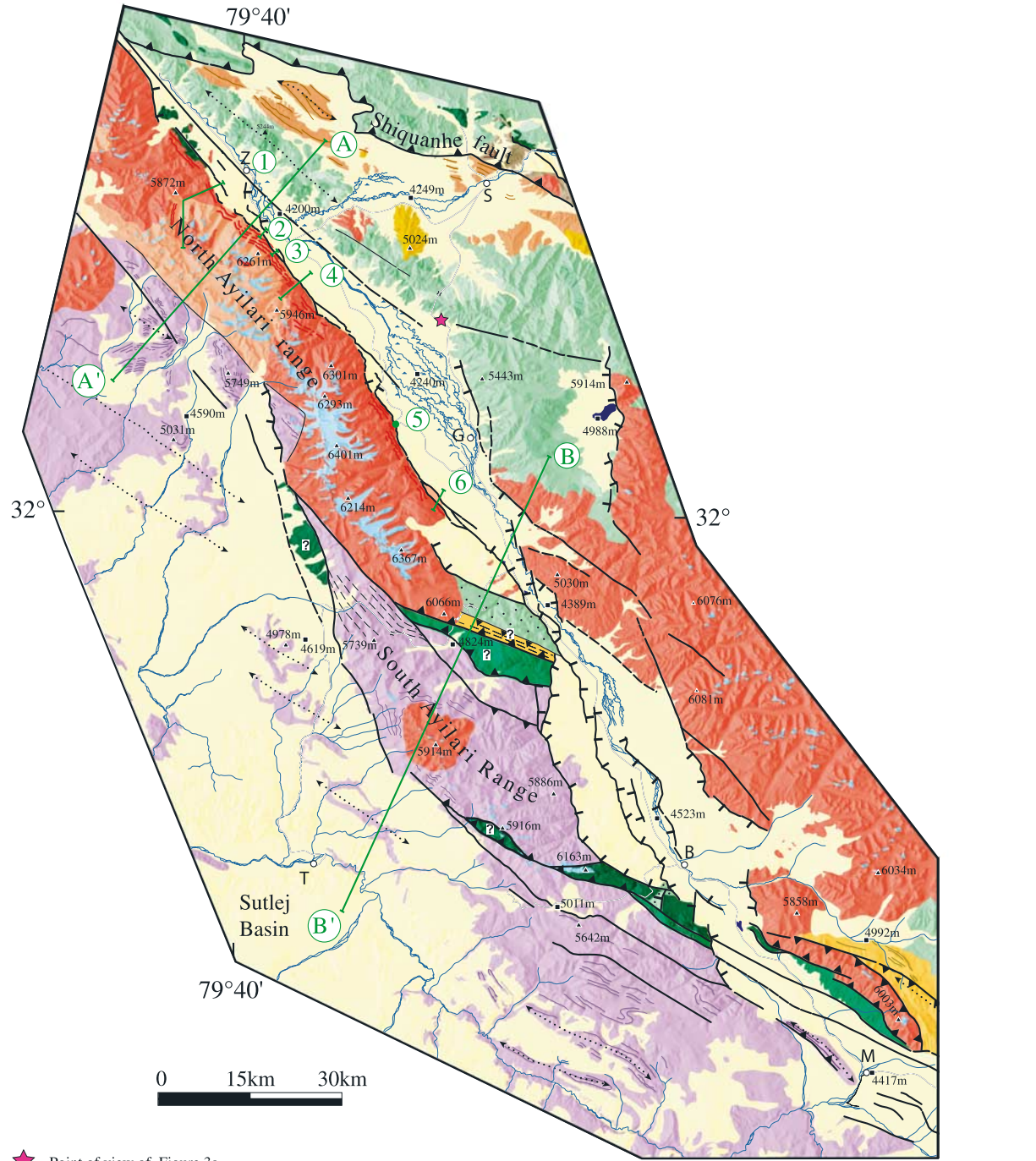


Figure 2

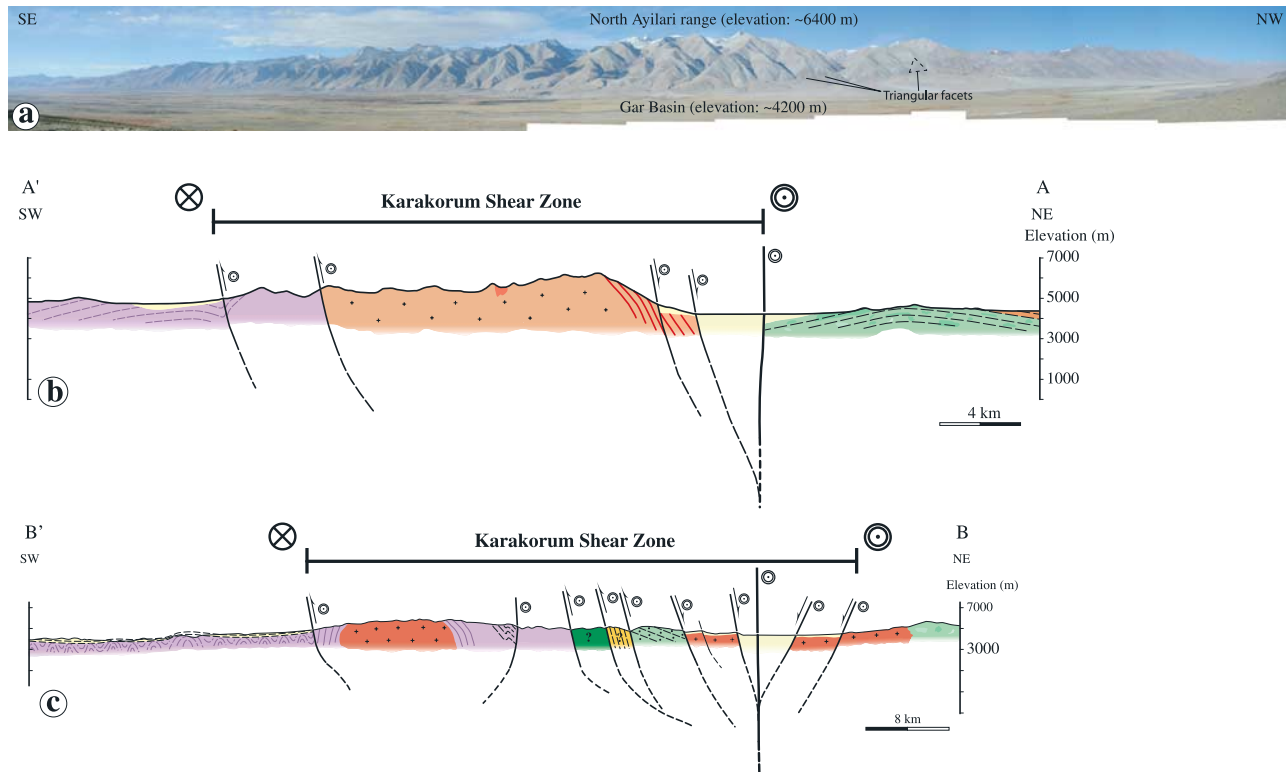


Figure 3. Panoramic field photograph of the Ayilari range and cross sections. Section locations and legend are on Figure 2. (a) Gar pull-apart basin and north Ayilari range seen from the northeastern side of the basin (point of view location reported on Figure 2). The main active dextral/normal strand of the KF follows the foot of the topography evidenced by cumulative several hundreds meters high triangular facets. (b) A-A' section across the north Ayilari range. (c) B-B' section passing through the southern Ayilari range. Some units were not clearly seen and are postulated by question mark. Movements on the southern strand of the KFZ are postulated from satellite imagery, and large-scale reconstruction is as postulated by *Lacassin et al.* [2004].

evolution of metamorphic rocks and allow estimates of the minimum long-term slip rate of the KFZ.

2. Geological Framework

[5] The KFZ is one of the most impressive morphological features of Asia, visible on all regional satellite images and digital elevation models (DEM). The active Karakorum fault is mainly strike slip and exhibits a 142°N trending, ~ 700 km long, linear trace, linking the Mustag Ata region in the Chinese Pamir ($\sim 38^{\circ}20'\text{N}$, $75^{\circ}10'\text{E}$) to the Shiquanhe area ($\sim 32^{\circ}30'\text{N}$, $80^{\circ}00'\text{E}$; Figure 1).

[6] South of Shiquanhe, the KFZ widens into a long (150×15 km) double rhomb-shaped basin limited by

NNW-SSE striking active faults (Figures 2 and 3) with geomorphic characteristics indicating a normal component of slip [*Armijo et al.*, 1989; *Lacassin et al.*, 2004; *Matte et al.*, 1996]. Southeastward, the active trace of the Karakorum fault (KF) continues for at least 150 km and becomes parallel to the Indus suture (105°N [*Lacassin et al.*, 2004; *Ratschbacher et al.*, 1994]).

[7] Along the southwestern edge of the northern rhomb-shaped basin (the Gar basin), cumulative scarps tens of meters high cutting late Pleistocene moraines, 2 km high triangular facets, and perched glacial valleys attest to a recent vertical throw on the fault (Figures 3 and 4) [*Matte et al.*, 1996]. These characters led *Armijo et al.* [1989] to interpret the Gar basin to be an active, dextral pull-apart

Figure 2. Geological map of the double rhomb-shaped Gar-Baer basin. Drawn from field observations, satellite image interpretation (Landsat 7, Spot 5, SRTM), and preexisting maps [*Armijo et al.*, 1989; *Jiao et al.*, 1988; *Kapp et al.*, 2003; *Matte et al.*, 1996; *Murphy et al.*, 2000]. Light shading outlines topography. Green lines show sections studied in the field (1 to 6, see Figure 4) and regional cross sections (A-A' and B-B' see Figures 3b and 3c). B, Baer; G, Gar; M, Menshi; S, Shiquanhe; T, Tsada; Z, Zhaxigang. Map is projected in UTM 44 using the ellipsoid WGS84.

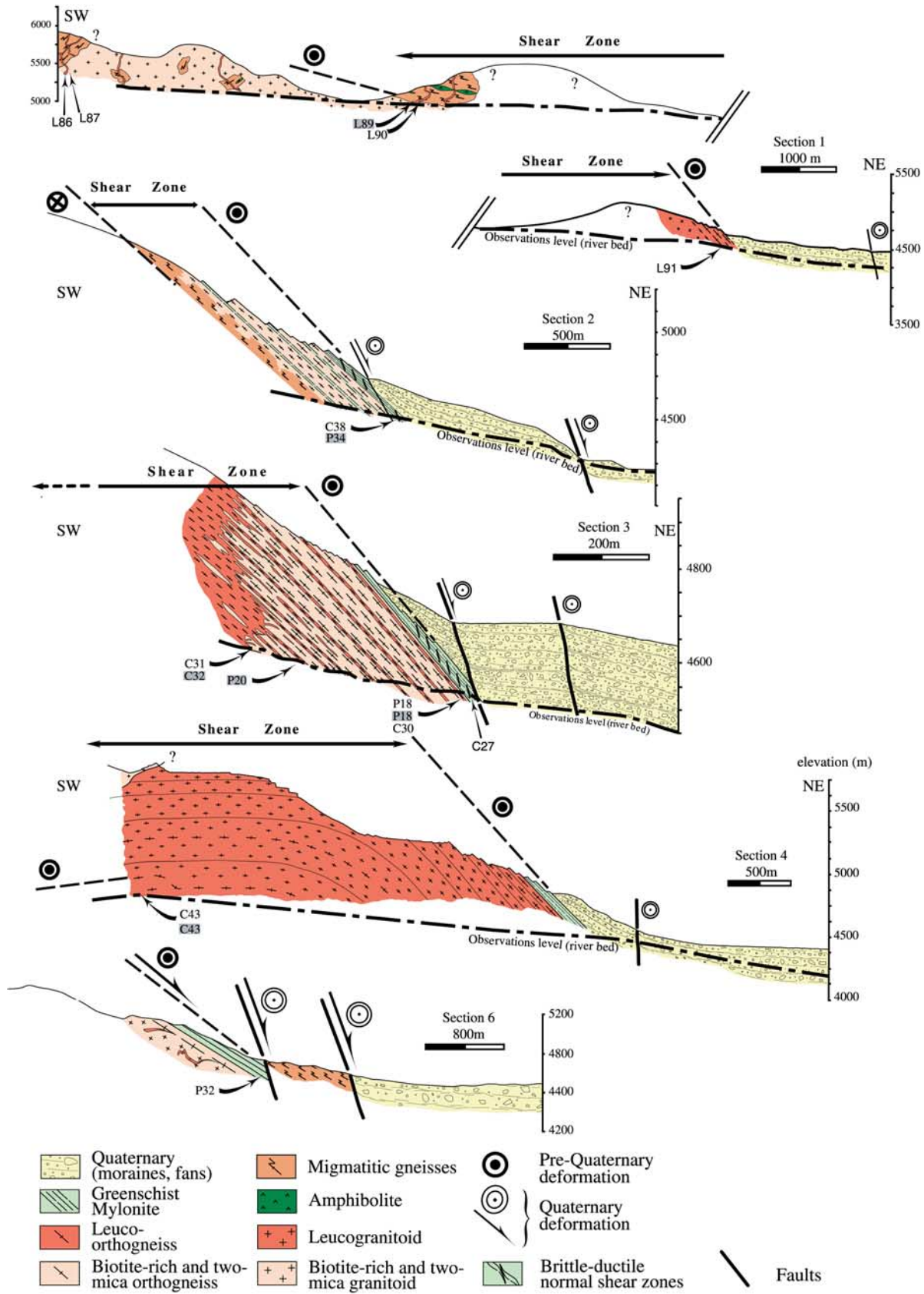


Figure 4

basin. Uplift due to movement along the active normal faults brings to the surface deformed metamorphic and magmatic rocks that form the North Ayilari range. The samples studied in this paper come from this range.

[8] Northwest of the Zhaxigang village (Z on Figure 2), and adjacent to the North Ayilari metamorphic and magmatic rocks, a sliver of ophiolitic rocks (gabbro, pyroxenite and basalt) outcrops between two strands of the active fault and most likely represents a piece of the Shyok suture dextrally offset by the Karakorum fault (Figures 1 and 2) [Lacassin *et al.*, 2004].

[9] Around Shiquanhe (S on Figure 2), Lower Cretaceous, shallow marine and terrestrial deposits, intruded by granitoids belonging to the Gangdese batholith [Schwab *et al.*, 2004], unconformably overlies kilometer wide massifs of ophiolitic mélange and locally Permian metasediments [Kapp *et al.*, 2003; Matte *et al.*, 1996]. These are remnants of a late Mesozoic suture [Cheng and Xu, 1987; Kapp *et al.*, 2003; Matte *et al.*, 1996; Ratschbacher *et al.*, 1994]. The Lower Cretaceous formation is, in turn, unconformably overlain by Tertiary red conglomerates [Matte *et al.*, 1996]. These series are cut by the Shiquanhe fault, which branches off the well-defined trace of the KF (Figure 2) [Ratschbacher *et al.*, 1994]. This fault bounds the Tertiary Shiquanhe basin to the north and thrusts the Permian-Tertiary sequence above vertical Pliocene-Quaternary sedimentary beds [Matte *et al.*, 1996]. Sedimentation in the Tertiary Shiquanhe basin was interpreted to be contemporaneous with displacement along the fault, suggesting an Oligocene onset of thrusting [Kapp *et al.*, 2003]. Crosscutting relationships and thermochronologic data show that the Shiquanhe fault is younger than the other thrusts parallel to it that are mapped northward up to the Risum anticlinorium (Jaggang, Zadoo, and Narangjiapo thrusts of Kapp *et al.* [2003, and references therein]). The thermal history calculated for a granite in the hanging wall of the Shiquanhe fault suggests a maximum shortening of 10 km across the fault since at least the Oligocene [Kapp *et al.*, 2003], whereas a minimum shortening of 3 km is estimated for the same period of time from the depth of the Tertiary Shiquanhe basin. The Shiquanhe fault may also have accommodated a recent dextral component of deformation and be connected to the N-S trending rifts of western Tibet [Armijo *et al.*, 1989; Ratschbacher *et al.*, 1994].

[10] South of the Ayilari range, the Sutlej basin is composed of conglomerates, sandstones, and siltstones cut by the impressive, up to 1000 m deep, canyons of the Sutlej river and its tributaries [Gansser, 1964]. The basement is composed of folded Meso-Cenozoic rocks pertaining to the Tethian series [Jiao *et al.*, 1988]. Fold axes have a roughly constant orientation of N125°E, oblique to the trend of the Ayilari range (Figure 2). Synsedimentary deformation features indicate that these folds were active during Mio-Pliocene time [Zhang *et al.*, 2000], and may still be active

since rivers form deep and narrow gorges where they cross anticlines. The contact between this basin and the Ayilari range is sharp, parallel to the active KFZ, and presents a strong topographic gradient (slope >25%) with a differential elevation of more than 1000m. This suggests recent vertical displacement along the southern flank of the Ayilari range. From large-scale mapping, Lacassin *et al.* [2004] postulated right-lateral strike-slip motion along this flank that would correspond to a southern strand of the KFZ (Figure 1).

[11] This southern fault strand would connect with the northern strand of the KFZ by ~N110°E trending faults, which isolate kilometer to tenths of kilometers wide slivers of Paleo-Cenozoic metasediments, black slates, meta graywackes, granitoids; ophiolites and conglomerates formed of peridotites, gabbros and basalts, a tectonic mélange consisting of blocks of serpentine, harzburgite, radiolarite, and Permo-Triassic hallstatt limestone, and Tertiary sandstone and conglomerate with boulders of volcanic rocks, granodiorite, sandstone and radiolarite (Figure 2). The slivers of ophiolites and of tectonic mélange probably represent the Indus suture zone dismembered by strike slip deformation [Armijo *et al.*, 1989; Lacassin *et al.*, 2004]. The Meso-Cenozoic meta graywackes might be deformed equivalents to the Cretaceous Xigaze group which widely outcrops ~400 km eastward [e.g., Burg, 1983] as a 20 km wide strip along the northern side of the Indus suture.

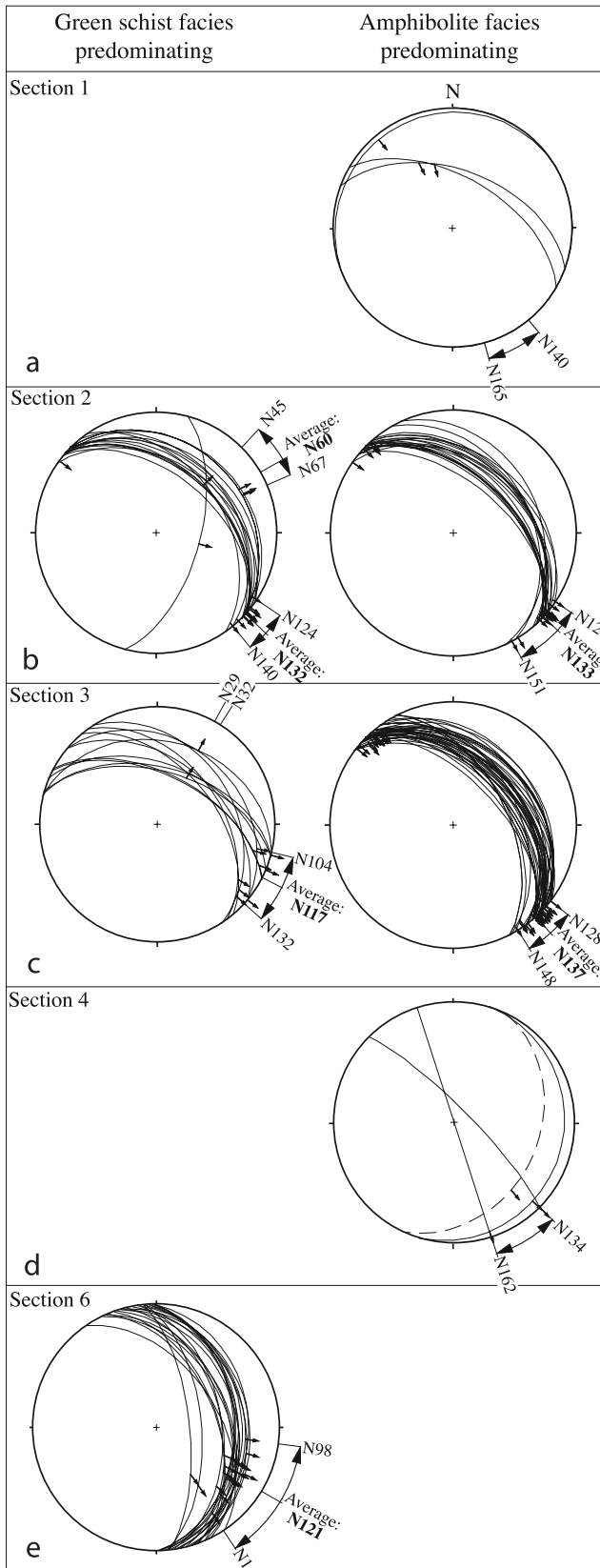
3. Field Observations in the North Ayilari Range

[12] The North Ayilari Range consists of high-grade metamorphic rocks and of variously deformed granitoids that we studied along six sections in valleys perpendicular to the trend of the range (Figures 2 and 4).

3.1. Structural Observations

[13] Frontal parts of sections 2, 3 and 6 show the development of strong cataclasis, including pseudotachylites, and low-temperature brittle-ductile deformation (Figure 4). The corresponding foliation was acquired in greenschist facies metamorphic conditions characterized by chlorite growth and is parallel to the active fault. The mean foliation trends ~N129°E and dips ~36°E on section 2, ~N123°E–45°NE for section 3, and ~N168°E–40°NE on section 6 (Figure 5). Lineations are underlined by chlorite and indicate a dextral-normal sense of shear characterized by either an oblique lineations with a pitch of 35–60°SE (section 6), or a partitioning between purely normal and purely strike slip deformation components with two groups of lineations trending N104–140°E and 29–67°NE, respectively (sections 2 and 3). This kinematics of low-grade deformation is consistent with the right-lateral motion deduced from the morphology along the adjacent active fault.

Figure 4. Five geological sections on the northern side of the north Ayilari range, west of Shiquanhe (location on Figure 2). Samples quoted were dated with Ar-Ar, (U-Th)/He, fission tracks (this study), or U-Th-Pb techniques (shaded samples [Valli, 2005; F. Valli *et al.*, manuscript in preparation, 2007]).



[14] This predominant greenschist facies foliation is superimposed upon, and partially overprints, a higher-grade deformation. A few tens of meters inside valleys 2, 3, and 6 and straight from the front of valleys 1, 4, and 5, chlorite is absent and the rocks were mainly deformed under higher-grade metamorphic conditions. Lineations, made of stretched feldspars and micas, are nearly horizontal and its trend is exceptionally constant from one section to another, and from the frontal parts of sections to ~10 km inside the range ($N139 \pm 6^\circ E$, at 2σ confidence level, Figures 5 and 6a). Foliation trend also remains roughly constant at $\sim N140^\circ E$; the dip angle decreasing progressively from the valley front ($60^\circ NE$, Figures 4 and 6b) to less than $10^\circ NE$, ~3–5 km inside the Ayilari range (sections 1 and 4). The sense of shear, mainly indicated by S-C structures (Figure 6c) and sigmoidal feldspars (Figure 6d), is unambiguously right-lateral where foliation is steep, and top to the south where it is gently dipping (section 1). Such geometry of the foliation and lineation is consistent with a convex Karakorum strike-slip shear zone with the rocks of the top northeastern edge displaced dextrally and southeastward.

3.2. Field Relationship, Samples Location, and Geochronological Framework

[15] In the following, the prefix “leuco” is used to qualify rocks rich in leucocratic minerals such as quartz-feldspar-muscovite (>95% of the rock volume). The attribute “two-mica” is assigned to rocks showing a content $\geq 5\%$ in both white and black mica, whereas the term “biotite-rich” applies to rocks having biotite and muscovite contents $>10\%$ and $<5\%$, respectively. In places biotite-rich, leucocratic and two-mica bands and lenses alternate, with the two systematically in contact. We call these rocks “migmatites,” as their biotite-rich, leucocratic and two-mica bands and lenses might represent mesosome, leucosome, and melanosome, respectively.

[16] Triangular facets in the northeastern part of section 1 (Figures 3a and 4) exhibit dextrally sheared leucogranitoids (sample L91). The intensity of deformation decreases progressively to a nondeformed leucogranitoid ~300 m from the front. Landslides deposited at high elevation by glacial moraines prevented us from mapping 5 km of the section, after which penetratively deformed migmatitic gneiss crops out (L89). There, the foliation is subhorizontal, the lineation strikes $\sim N153^\circ E$ roughly parallel to the KF, and the sense of shear, indicated by S-C structures and sigmoidal feldspars, is top to the south. The foliation is locally crosscut by leucocratic veins (L86, L90) that are differentiated magmas originating from a large two-mica granitoid body located ~500 m southwestward (L87). U-Th-Pb dating on zircons

Figure 5. Structural data from the North Ayilari range (lower hemisphere Schmidt diagrams). Great circles are foliation planes; dashed great circles represent C (shear) planes; and arrows show the direction of lineations with the arrow pointing toward the hanging wall motion. Extreme and average values of lineation strikes are indicated. (left) Greenschist and (right) higher-grade metamorphic facies.

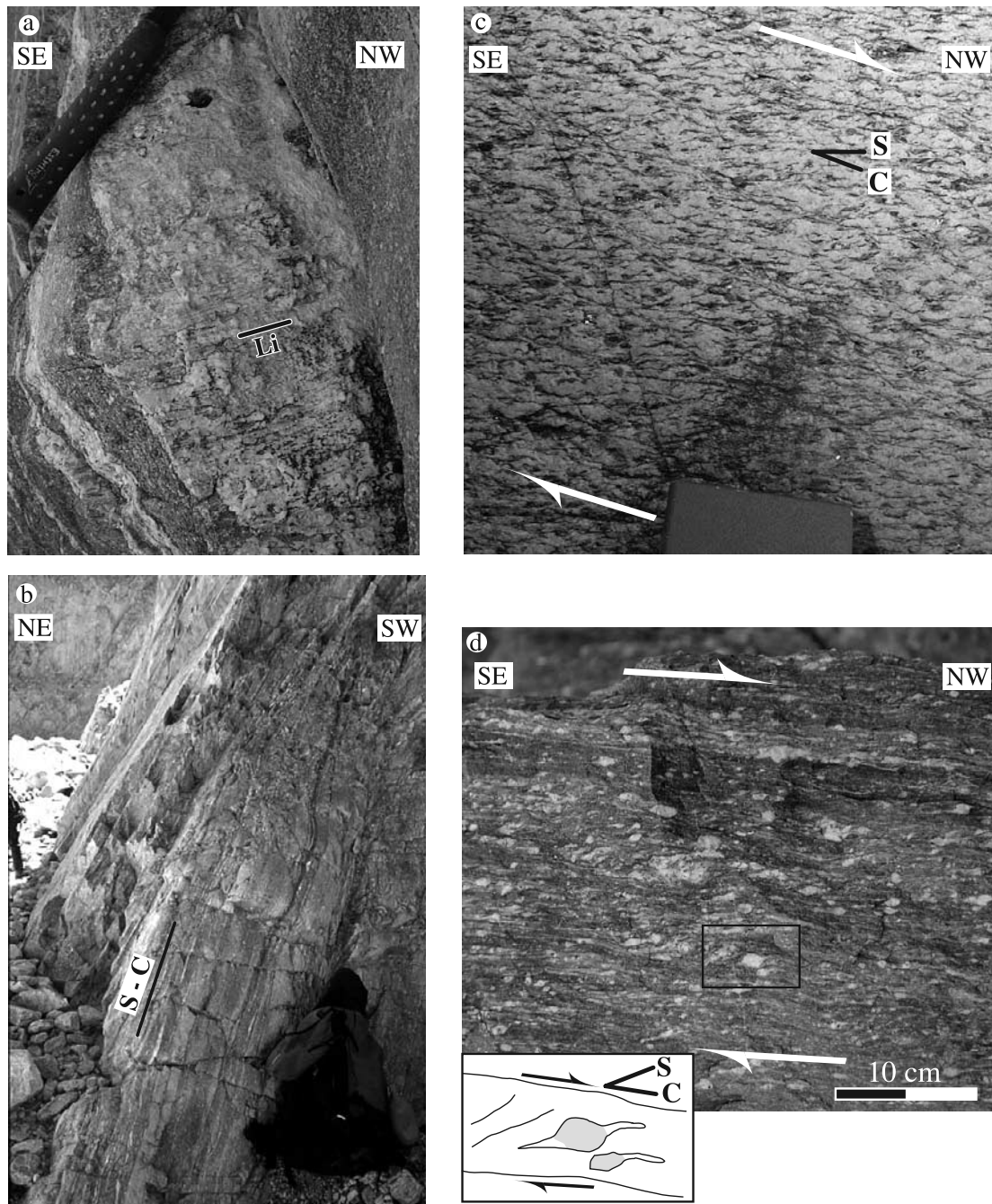


Figure 6. Characteristics of deformation along the Karakorum shear zone. (a) Steep foliation with an horizontal lineation made of stretched feldspar, quartz, and mica in a leucocratic vein (section 2, Figure 2). The hammer is for scale. (b) Steep foliation (S) transposed along the shear (C) planes in two-mica gneisses close to the frontal part of valley (section 3, Figure 2). The bag gives scale. (c) Dextrally sheared leucogranitoid seen from above (section 3, Figure 2). The compass is for scale. (d) Two-mica gneiss exhibiting dextrally sheared sigmoidal feldspars (section 2, Figure 2). View from above.

of sample L89 suggest crystallization of the migmatitic gneiss at 23.4 ± 4.0 Ma, while monazites revealed post crystallization event(s) interpreted as hydrothermal activity until ~ 18 Ma (L89 [Valli, 2005; F. Valli et al., Ductile deformation and magmatism along the Karakorum fault,

western Tibet: U-Th-Pb geochronological constraints, manuscript in preparation, 2007, hereinafter referred to as F. Valli et al., manuscript in preparation, 2007]).

[17] On section 2 (Figure 4), two-mica gneisses, mylonites, and ultramylonites crop out as a band at least 500 m

Table 1. Mineralogical Composition and Grain Size of the Dated Samples^a

Samples	Minerals										Minerals Dated	Temperature of Deformation, °C	
	Quartz	Feldspar	Mus	Bio	Chl	Grt	Tit	Ox	Ep	Ap			
C27	25%, <0.4	35%, <1.5	10%, <0.01	-	10%	-	-	10%	10%	-	-	Kfs	from >600-400 to <250
C30	40%, <0.5	35%, <5	20%, <5	-	1%	-	-	<1%	<1%	-	-	Kfs, Ms	from >600-400 to <250
C31	50%, <0.5	40%, <1.5	10%, <2	<1%	<1%	5%, <1	-	-	-	-	-	Kfs, Ms	from >600-400 to <250
C38	50%, <0.5	35%, <2	15%, <1	-	<1%	-	-	-	-	-	-	Kfs, Ms	from >600-400 to <250
C43	30%, 0.1-1	55%, 0.1-1.5	10%, <5	3%, <5	2%	-	-	<1%	-	-	-	Kfs, Ms, Bt	from >600-400 to <250
L86	35%, ~1	40%, 1-5	20%, ~1	5%, ~1	<1%	-	-	<1%	-	-	-	Ms	no deformation
L87	30%, <0.5	45%, 0.5-10	10%, ~0.5	15%, ~0.5	<1%	-	-	<1%	-	-	-	Ap	no deformation
L90	40%, ~1	35%, 1-5	15%, ~1	5%, ~1	<1%	-	-	<1%	-	-	-	Kfs, Ms	from >600-400 to <250
L91	40%, <0.5	50%, <1.5	8%, <1	1%	1%	-	-	<1%	-	-	-	Kfs, Ms, Bt	from >600-400 to <250
P18	20%, <0.1	40%, <10	20%, <1	15%, <1	-	-	-	<1%	-	-	-	Ap	from >600-400 to <250
P32	30%, <0.5	50%, <10	10%, <0.01	5%	10%	1%	-	3%	1%	<1%	-	Kfs, Bt, Ap	from >600-400 to <250
P56	30%, <1	45%, <3	5%, <0.01	15%	5%	-	-	<1%	-	-	-	Bt	no deformation
P57	30%, <1	45%, <3	5%, <0.01	15%	5%	-	-	<1%	-	-	-	Kfs, Bt	no deformation

^aPercentages are estimated from thin section with an optical microscope; crystal sizes are quoted in centimeter. Abbreviations are from *Kretz* [1983] (Tit, titanite).

thick. Inside the valley, migmatitic gneisses crop out. The leucocratic mylonite sample C38 from the frontal greenschist part was selected for Ar/Ar dating, whereas a similar sample (P34) yielded metamorphic and hydrothermal zircons 22.0 ± 4.7 Ma old [*Valli, 2005*; F. Valli et al., manuscript in preparation, 2007].

[18] Frontal greenschist gneisses (C27) of section 3 (Figure 4) pass southwestward to high-grade biotite-rich and two-mica gneisses (P18), and mylonites, interlayered with sheared leucocratic veins (C30). Leucocratic veins increase in frequency toward the southwest, where leucogranitic pods (C31) invade biotite-rich and two-mica gneisses and lead to a complex set of deformed dikes which are connected to a dextrally sheared leucogranitoid massif (C32). The leucocratic veins are more or less deformed, with the oldest ones transposed in the foliation of the hosting gneisses and highly sheared [see *Lacassin et al., 2004, Figure 4d*]. Strain intensity in the veins increases with decreasing angles between the vein and the foliation in the hosting gneisses. The younger veins crosscut the hosting gneisses foliation with angles up to 90° [*Lacassin et al., 2004, Figure 4c*] but still exhibit a slight foliation parallel to the one of the hosting rocks. This indicates that veins emplaced successively during right-lateral shear have been more or less rotated and deformed depending on their age [*Lacassin et al., 2004; Valli, 2005*].

[19] A concordant fraction of needle-like zircon grains yield the ~ 23 Ma crystallization age of the synkinematic leucogranitoid C32 [*Lacassin et al., 2004; Valli, 2005*]. Zircons and monazites from the hosting biotite-rich and two-mica gneisses indicate that they are ortho-derived and crystallized between ~ 26.8 and 19.2 Ma (samples P18 and P20 [*Valli, 2005*; F. Valli et al., manuscript in preparation, 2007]). The crystallization of monazites occurred between ~ 19 Ma and ~ 16 Ma, likely during hydrothermal alteration (samples P18 and C32).

[20] Section 4 (Figure 4) is exclusively composed of sheared leucogranitoid. A deformed sample was taken at the southwest end of the section (C43). U-Th-Pb dating of monazites and zircons revealed the crystallization of hydrothermal and metamorphic zircons and monazites occurred between ~ 14 and ~ 22 Ma.

[21] Outcrop 5 (Figure 2) corresponds to a nondeformed two-mica granitoid (P56, P57).

[22] Section 6 (Figure 4) exhibits intensively fractured migmatites, dextral-normal greenschist two-mica gneisses (P32), and finally a slightly deformed two-mica granitoid crosscut by variously deformed tourmaline-bearing leucocratic veins.

4. Petrostructural Evidence for Continuum Shearing From High to Near-Surface Temperatures

[23] Rocks are mainly composed of quartz, feldspar and mica; other constituents, proportions and crystal sizes are detailed in Table 1. Samples P56, P57, L86, and L87 are isotropic and present a plutonic texture, they are herein

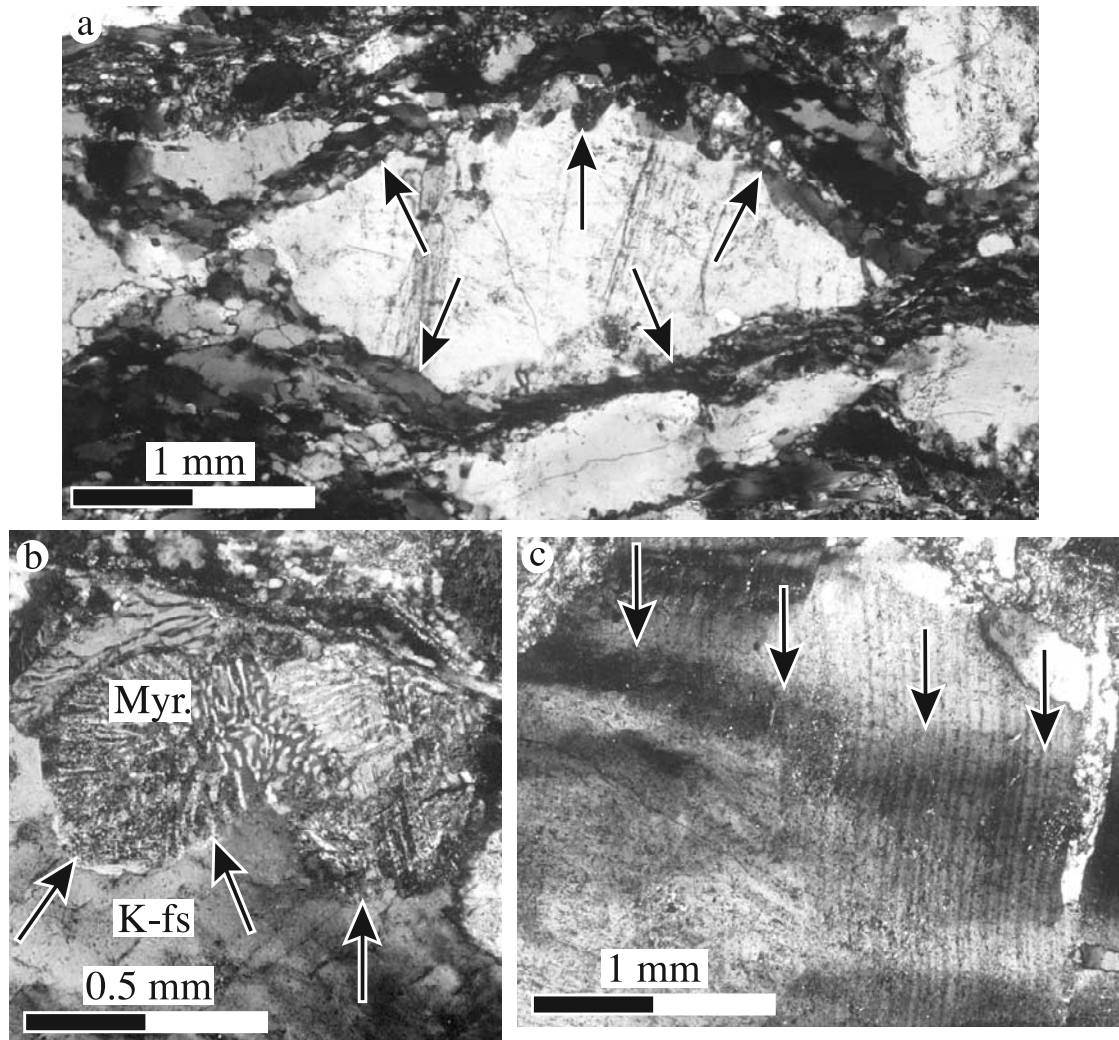


Figure 7. Optical photomicrographs under cross-polarized light of feldspar microstructures. (a) Feldspar porphyroblast surrounded by a mantle of recrystallized feldspars (arrows, sample C38). (b) Well-developed myrmekite (Myr) on a K-feldspar grain (K-fs) boundary (arrows, sample P32). (c) Bending of feldspar twin lamellae by dislocation glide (arrows, sample C29).

called granitoids. All other samples (L90, L91, C27, C30, C31, C38, C43, P18, P32) are deformed to various extents under various temperature conditions.

[24] Feldspar porphyroclasts commonly show recrystallized grains at their boundaries, producing a core-and-rim structure diagnostic of dynamic recrystallization [e.g., *Passchier and Trouw, 1996*] (Figure 7a). Such dynamic recrystallization strongly depends upon temperature, and to a minor extent on other factors such as strain rate, differential stress, and the chemical activity of certain components such as water [*Passchier and Trouw, 1996; White, 1975*]. Consistent observations in several natural examples suggest that such dynamic recrystallization occurs at medium- to high-temperature conditions (>600–400°C) during deformation [e.g., *Gapais, 1989a, 1989b; Leloup et al.,*

1995; Passchier and Trouw, 1996]. Other deformation characteristics in feldspar indicate that deformation continued under retrograde conditions. Abundant myrmekite growth along the boundaries of K-feldspar porphyroclasts (Figure 7b), especially at high stress sites, suggests a deformation under medium-grade temperature conditions (400–600°C [*Borges and White, 1980; Gapais, 1989a, 1989b; Gates and Glover, 1989; Passchier and Trouw, 1996; Simpson and Wintsch, 1989; Tsurumi et al., 2003; Tullis and Yund, 1991*]). Bent twin lamellae (Figure 7c), undulose extinction, and kink bands suggest deformation under lower-grade temperature conditions (300–400°C [e.g., *Gower and Simpson, 1992; Jensen and Starkey, 1985; Ji and Mainprice, 1990; Olesen, 1987; Olsen and Kohlstedt, 1985; White and Mawer, 1986*]). Finally, brittle

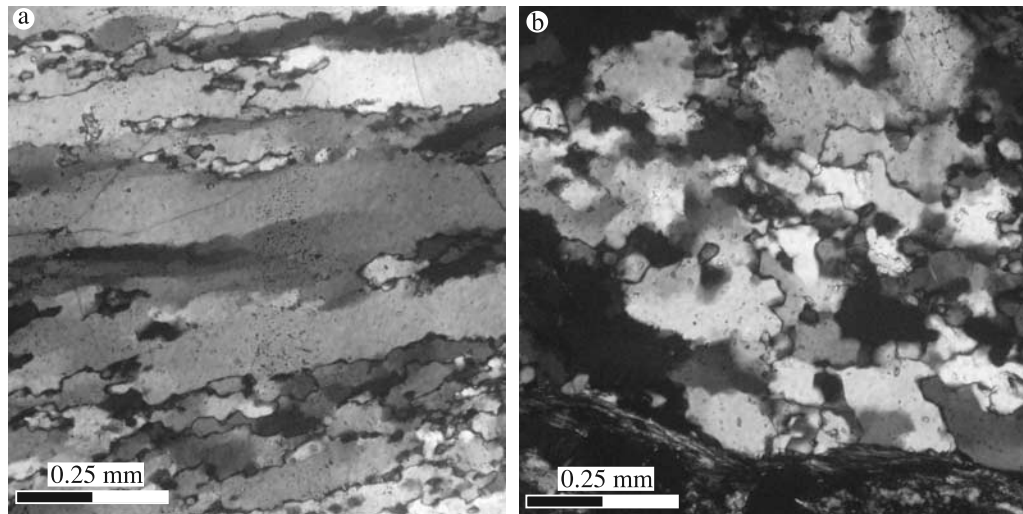


Figure 8. Optical photomicrographs under cross polarized light of quartz microstructures. (a) Relics of large old quartz grains with undulose extinction and elongate subgrains passing laterally into domains of small grains dynamically recrystallized, probably by subgrain rotation recrystallization (sample P32). (b) Polycrystalline quartz with irregular grain boundaries formed in response to grain boundary migration recrystallization (sample P32).

deformation of porphyroclasts, often in the form of “book-shelf” microfracturing (samples C30, P32), and, in place, patchy undulose extinction indicates even lower-grade conditions (below 300°C [Passchier, 1982; Passchier and Trouw, 1996; Tullis and Yund, 1987]) during late increments of deformation.

[25] Dynamic recrystallization through subgrain rotation (Figure 8a) or grain boundary migration (Figure 8b) occurs in quartz [Guillopé and Poirier, 1979; Passchier and Trouw, 1996; Tullis and Yund, 1991]. These structures are typical at medium- to high-grade conditions (700–400°C [Hirth and Tullis, 1992; Passchier and Trouw, 1996]). Quartz grains also exhibit undulose extinctions typical of low-grade conditions below 300°C [Passchier and Trouw, 1996] during late deformation stages.

[26] Mica fish are thought to form due to combined slip on the basal plane, rigid body rotation, boudinage and recrystallization at the edges [Eisbacher, 1970]. These mechanisms imply deformation at temperature higher than ~250°C [Stesky, 1978; Stesky et al., 1974]. Micas also show characters typical of very low-grade metamorphism below ~250°C, such as undulose extinction and, in some samples, kinking and folding (P36, C43, [Stesky, 1978; Stesky et al., 1974]).

[27] The different temperature-dependent deformation mechanisms described above produce asymmetric microstructures indicating the sense of shear. Sigmoidal and δ -type feldspars, quartz subgrains rotation and recrystallization, myrmekite formation at high stress sites, and asymmetric mica “fishes” all indicate the same unambiguous dextral shear sense on NE dipping foliations, or top to the south shear in samples with flatter foliation planes. Such dextral

and top to the south shear thus appears to have been continuous from temperature higher than 400–600°C down to low-grade conditions of <250°C.

5. Analytical Methods

5.1. The ^{40}Ar - ^{39}Ar Techniques

[28] Minerals were separated using heavy liquids, a Frantz magnetic separator and finally by hand picking under a binocular microscope. The samples were irradiated at the Phoenix Memorial Laboratory reactor of the University of Michigan, in the L67 position for 20 h with a 10^{18} neutrons $\text{cm}^{-2} \text{s}^{-1}$ flux. Irradiation interference on K, Ca, and Cl were corrected by irradiating and analyzing KCl and CaF_2 pure salts. J factors were estimated by the use of duplicates of the Fish Canyon sanidine standard with an age of 28.02 Ma [Renne et al., 1998].

[29] The samples were analyzed in Clermont-Ferrand (samples L90, L91, L86, C38, C27, C29, C30, C31, C43) and Montpellier (samples P32, P56 and P57) using the same apparatus and the same protocol, as described by Arnaud et al. [2003]. Samples were loaded in aluminum packets into a double vacuum Staudacher-type furnace and step heated; temperature is calibrated by means of a thermocouple. The gas was purified using cold traps with liquid air and Al-Zr getters. Once cleaned, the gas was introduced into a VG3600 mass spectrometer and allowed to equilibrate for 2 min prior to analysis was done statically. Signals were measured by the mean of a Faraday cup with a 10^{11} ohm resistor for ^{40}Ar and ^{39}Ar while ^{39}Ar , ^{38}Ar , ^{37}Ar , and ^{36}Ar were analyzed with a photomultiplier after interaction on a

Table 2. Apatite (U-Th)/He Ages^a

	Location	Mass, μg	Mean Radius, μm	Mean Length, μm	U, ppm	Th, ppm	Raw Age, Ma	He, nmol g^{-1}	Ft	Corrected Age, Ma	Uncertainty 1σ
K3P32A	N32°01'38.4, E80°00'38.2	41.1	81.4	282.8	10	24.04	3.1	0.26	0.82	3.7	0.1 ^b
K3P32B	N32°01'38.4, E80°00'38.2	34.6	69.3	336.4	12.91	32.48	3.7	0.42	0.8	4.7	0.1 ^b
K3P32C	N32°01'38.4, E80°00'38.2	17.2	56.4	250.7	14.43	35.49	3.6	0.45	0.75	4.8	0.2 ^b
Mean										4.4	0.4

^aAges corrected for α ejection following method of *Farley et al.* [1996]. Analytical procedure after *House et al.* [2000].

^bAnalytical uncertainty.

Daly plate. Gain between both collectors was estimated by duplicate analysis of ^{39}Ar on both collectors during each analysis, and also by statistical analysis over a period of several years. This gain is 95 and is known at better than 1.5%. This error is included in the age calculation, along with analytical errors on each signal and errors on the blank values. Detailed analytical results are available in the auxiliary material.¹ Age plateau given are weighted mean plateaus; the error takes the error on the J factor into account. With the historical decrease of analytical errors, strict plateau criteria [*Berger and York*, 1981; *Dalrymple and Lanphere*, 1974] are less frequently met. Thus pseudo-plateaus are used when a significant number of steps overlap globally at 2σ even if contiguous steps do not. For K-feldspars, plateau ages cannot be defined, but since we wish to compare and discuss series of steps with similar ages we used simple mean, thus unweighted, ages. Isochron ages are obtained on an inverse isochron diagram of $^{36}\text{Ar}/^{40}\text{Ar}$ versus $^{39}\text{Ar}/^{40}\text{Ar}$ [*Roddick*, 1978; *Roddick et al.*, 1980], which often allows homogeneous excess components to be identified. Errors on age and intercept age include individual errors on each point and linear regression by *York's* [1969] method. The goodness of fit relative to individual errors is measured by mean square weighted deviation (MSWD).

[30] For micas, classical furnace step heating was conducted and usually yielded an almost perfectly flat age spectra, from which plateau and isochron ages were calculated and are shown side by side to assess potential excess argon problems. If the inverse isochron age is close to the plateau age and $^{40}\text{Ar}/^{36}\text{Ar}$ is not significantly different from present-day $^{40}\text{Ar}/^{36}\text{Ar}$ atmospheric ratio (295.5), we consider that the plateau age is reliable. When this is not the case, we suspect a nonatmospheric initial $^{40}\text{Ar}/^{36}\text{Ar}$ ratio and we thus prefer to rely on the inverse isochron age if this one is well determined.

[31] A more peculiar step-heating schedule was conducted on K-feldspar in order to retrieve diffusion characteristics, to apply diffusion models, and to calculate model thermal histories [*Harrison et al.*, 1991; *Lovera et al.*, 1989, 1991]. We conducted duplicated isothermal step heating at

low temperatures (450–800°C); these often yield a saw tooth-shaped age spectrum where the second of the two stages is systematically younger and probably less affected by excess argon [e.g., *Harrison et al.*, 1994]. The use of pseudoplateaus or isochrons on series of steps on K-feldspars may seem at odds with the classical definitions. However, K-feldspars are known to be composite and also to record, in certain conditions, cooling histories and especially rapid cooling events, which are theoretically, marked by flat portions of the age spectra [*Lovera et al.*, 1989, 1991]. Those portions can be treated by statistical approaches (such as plateaus and isochrons) and can be interpreted as revealing undisturbed quenched parts of the sample or alternatively the presence of excess argon. All errors are quoted at 2 sigmas.

5.2. (U-Th)/He Techniques

[32] We used euhedral apatite grains that were checked for inclusions under a binocular microscope. Grain dimensions were measured under a microscope to determine the α emission correction [*Farley et al.*, 1996]. All replicates were loaded in platinum capsules and heated for He extraction by laser for 5 min at $\sim 1050^\circ\text{C}$ following the procedures described by *House et al.* [2000]. Packets were then transferred into Teflon beakers for dissolution before being analyzed with a Finnegan element ICP-MS (see *House et al.* [2000] for detailed analytical procedure). The mean age is reported as there were three replicate analyses (Table 2). Propagated errors on He ages based on the analytical uncertainty in U, Th, and He measurements are 2% (1-sigma) for laser samples [*Farley*, 2002]. However, the propagated error from geometrical correction of ages as well as intrinsic heterogeneities in U content usually lead to much higher errors. *Farley* [2002] suggests a minimum uncertainty of 3% (1-sigma) for all ages, based on the reproducibility of laboratory standards. Mean error is reported at 1-sigma as standard errors using the standard deviation of the replicate analyses divided by $(n-1)^{1/2}$ where n is the number of replicate analyses performed (Table 2). This error estimate is larger than the analytical error alone and likely reflects the age uncertainty due to differences in grain size, zoning of parent material and other factors which may contribute to grain differences in He age and uncer-

¹Auxiliary materials are available in the HTML. doi:10.1029/2005TC001913.

Table 3. K-Feldspar $^{40}\text{Ar}/^{39}\text{Ar}$ Results^a

Section/Site	Number	Location	Statistics	Plateau Age			Inverse Isochron Age				Inverse Isochron Age on All Steps			
				Age, Ma	Steps	Age, Ma	$^{40}\text{Ar}/^{36}\text{Ar}_i$	MSWD	Steps	Age, Ma	Steps	Age, Ma	$^{40}\text{Ar}/^{36}\text{Ar}_i$	MSWD
Section 1/frontal part	L90	32°23'45.8"N 79°33'30.0"E	Plateau	LT: 13.2 ± 0.4	4 steps/26 (2-5)	12.7 ± 0.2	568 ± 25	0.6	14 steps/26 (2-15)	24 steps/26 (2-25)	12.6 ± 0.2	593 ± 12	2.36	
			SMA	HT: 21.2 ± 1.0	10% of gas 13 steps/26 (13-26)	-	-	-	31% of gas	-	-	-	-	-
Section 1/frontal part	L91	32°28'46.9"N 79°36'45.0"E	Plateau	LT: 12.1 ± 0.3	8 steps/26 (2-9)	11.9 ± 0.4	719 ± 120	0.6	14 steps/26 (2-15)	24 steps/26 (2-25)	11.9 ± 0.2	857 ± 79	4.5	
			SMA	HT: 15.7 ± 0.4	68% of gas 4 steps/26 (21-25)	-	-	-	34% of gas	-	-	-	-	-
Section 2/frontal part	C38	32°25'01.1"N 79°42'09.4"E	Plateau	LT: 9.3 ± 0.3	4 steps/25 (2-5)	9.1 ± 0.4	540 ± 80	4.1	12 steps/25 (2-13)	26 steps/26	8.2 ± 0.4	934 ± 73	15.5	
			SMA	HT: 18.4 ± 1.3	9% of gas 4 steps/25 (22-25)	-	-	-	21% of gas 4 steps/25 (22-25)	-	-	-	-	-
Section 3/frontal part	C27	32°23'28.9"N 79°43'35.7"E	Plateau	LT: 9.0 ± 0.3	4 steps/26 (2-5)	8.8 ± 0.4	533 ± 34	3.5	13 steps/26 (2-14)	25 steps/25	8.1 ± 0.5	879 ± 84	25	
			SMA	HT: 15.3 ± 0.6	44% of gas 10% of gas	-	-	-	24% of gas	-	-	-	-	-
Section 3/frontal part	C30	32°23'28.9"N 79°43'35.0"E	HT at highest step	20.8 ± 0.3	-	-	-	-	-	-	-	-	-	
			Plateau	LT: 9.0 ± 0.3	4 steps/26 (2-5)	8.7 ± 0.3	600 ± 54	1.9	12 steps/26 (2-13)	26 steps/26	8.0 ± 0.4	806 ± 55	9.3	
Section 3/~500 m from front	C31	32°23'17.6"N 79°43'25.3"E	Plateau	HT: 15.3 ± 0.6	7% of gas 3 steps/26 (23-25)	-	-	-	18% of gas	-	-	-	very bad	
			Plateau	LT: 9.7 ± 0.2	44% of gas 6 steps/26 (2-7)	8.8 ± 0.8	807 ± 200	10	14 steps/26 (2-15)	26 steps/26	7.6 ± 0.7	1220 ± 163	30	
Section 4/~3.5 km from front	C43	32°19'34.9"N 79°44'21.1"E	SMA	HT: 18.7 ± 0.4	15% of gas 3 steps/26 (23-25)	-	-	-	30% of gas	-	-	-	-	
			Plateau	LT: 10.1 ± 0.2	37% of gas 4 steps/26 (2-5)	9.8 ± 0.4	606 ± 44	2.1	14 steps/26 (2-15)	24 steps/26 (2-26)	9.8 ± 0.3	632 ± 33	12	
Outcrop 5	P56	32°08'14.7"N 79°56'12.9"E	Plateau	HT: 16.3 ± 0.6	11% of gas 3 steps/26 (24-26)	-	-	-	29% of gas	-	-	-	-	
			SMA	LT: 7.3 ± 1.0	35% of gas 11 steps/27 (4-14)	-	-	-	-	-	-	-	-	-
			SMA	HT: ≥ 16	18% of gas	-	-	-	-	-	-	-	-	

Table 3. (continued)

Section/Site	Sample		Plateau Age		Inverse Isochron Age			Inverse Isochron Age on All Steps				
	Number	Location	Statistics	Age, Ma	Steps	Age, Ma	$^{40}\text{Ar}/^{36}\text{Ar}_i$	MSWD	Steps	Age, Ma	$^{40}\text{Ar}/^{36}\text{Ar}_i$	MSWD
Outcrop 5	P57	32°08'14.7"N 79°56'12.9"E	SMA	LT: 8.3 ± 1.0	6 steps/31 (3–8) 9% of gas	8.0 ± 1.0	358 ± 60	30	18 steps/31 (2–19) 35% of gas	-	-	-
Section 6/frontal part	P32	32°01'38.4"N 80°00'38.2"E	SMA Plateau	HT: ≥ 16 LT: 9.0 ± 1.4	4 steps/30 (2–5) 6% of gas	8.7 ± 0.8	295 ± 29	2.8	4 steps/30 (2–5) 6% of gas	-	-	-

^aLT, low furnace temperature (400–700°C) age; HT, high furnace temperature (1000–1400°C); SMA, simple mean age. Errors are quoted at 2σ confidence level.

tainties in the alpha ejection correction [i.e., House *et al.*, 2001].

5.3. Apatite Fission Track Methodology

[33] Apatite was separated using standard magnetic and density methods, then mounted on glass slides with araldite epoxy. After grinding and polishing to expose an internal surface, the apatites were etched with 5.5 molar nitric acid for 20 seconds at 21°C. Samples were irradiated at Oregon State University. After irradiation, mica external detectors were etched in 40% HF for 45 min at 21°C. For age determinations, 20 to 23 good quality grains per sample were selected at random and dated using the external detector method. Laboratory procedures are essentially the same as those reported by Sobel and Strecker [2003].

[34] At temperatures hotter than ~110–150°C, all fission tracks are totally annealed, resetting the fission track clock to zero. The total annealing temperature (T_A) and the effective closure temperature depend on the kinetic characteristics of the apatite and the cooling rate [Ketcham *et al.*, 1999]. The partial annealing zone (PAZ) extends from the T_A down to ~60°C, and within this temperature range tracks are partially annealed. Below ~60°C, apatite fission track (AFT) is effectively stable because annealing occurs at a very slow rate [e.g., Gleadow *et al.*, 1986]. Following convention, all statistical uncertainties on pooled ages and mean track lengths are quoted at the ±1σ level, but ±2σ uncertainties are taken into account for geologic interpretation.

6. Analytical Results

6.1. K-Feldspar

[35] The shapes of all age spectra are roughly similar, with a first low-temperature (LT) pseudoplateau between ~10 and 40% of gas release corresponding to furnace temperature below 700°C, followed by a rapid age increase toward a high temperature (HT) plateau or pseudoplateau (Table 3, Figure 9a, and auxiliary material Figures S1–S10).

[36] The youngest steps on K-feldspar LT ages (L90, L91, C27, C30, C38, C43, P32, P57), or pseudoplateaus (C31, P56) defined on K-feldspar LT steps, appear diachronous between the northernmost section and other localities: two samples from section 1 yield ages at ~14–12 Ma, whereas in other sections (sections 2, 3, 4, 5, and 6), the eight K-feldspar samples invariably exhibit LT pseudoplateau age in a range of ~10–7 Ma. Inverse isochron diagrams on the same LT steps often revealed slight excess argon, in agreement with most studies revealing excess argon release from K-feldspars at low temperature due to fluid inclusion breakdown [Burgess *et al.*, 1992; Harrison *et al.*, 1994]. In our case, excess argon had only a minor effect since the isochron age estimates are similar to the pseudoplateau ones when considered at the 2σ confidence level.

[37] The HT K-feldspar age plateau (L90, L91, C30, C31, C38, C43), or pseudoplateau (C27), yield a relatively wide range of average ages from 21.2 ± 1.0 to 15.3 ± 0.7 Ma (Table 3), with no correlation with the geographic position

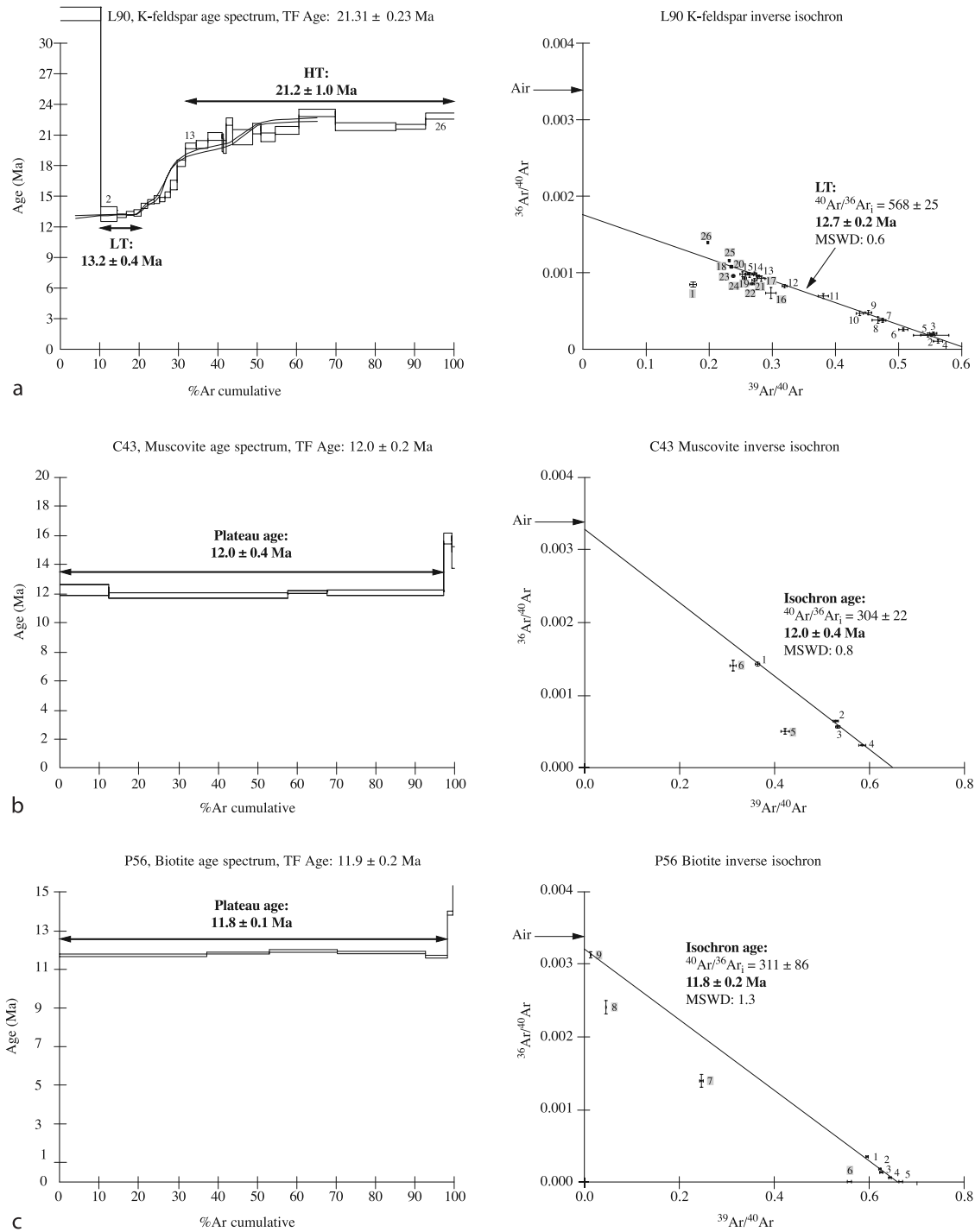


Figure 9. Typical $^{40}\text{Ar}/^{39}\text{Ar}$ results from the north Ayilari range. Data are summarized in Table 3, 4, and 5 and samples are located on Figure 4. All other samples are shown in the auxiliary material. (left) Conventional age spectra. Double arrows represent age plateaus or pseudoplateaus (see text for explanation on the calculation and use of these terms). (right) Inverse isochrones for which each step is plotted with the corresponding error bars and step number. The regression line corresponds to the calculated age and $^{40}\text{Ar}/^{36}\text{Ar}_i$. Step numbers in shaded boxes were not used for calculations. (a) K-feldspar (sample L90). Superimposed curves represent synthetic age spectra resulting from thermal modeling at 2σ confidence level. (b) Muscovite (sample C43). (c) Biotite (sample P56). Data from other samples, and K/Ca, Cl/K plots are reported in the auxiliary material: Figures S1–S10 (K-feldspar), S11–S17 (muscovite), and S18–S22 (biotite).

Table 4. Muscovite $^{40}\text{Ar}/^{39}\text{Ar}$ Results^a

Section/Site	Sample		Plateau Age			Inverse Isochron Age			Total Fusion Age, Ma	
	Number	Location	Statistics	Age, Ma	Steps	Age, Ma	$^{40}\text{Ar}/^{36}\text{Ar}_i$	MSWD		Steps
Section 1/frontal part	L91	32°28'46.9"N 79°36'45.0"E	Plateau	14.7 ± 0.1	4 steps/8 (1–4) 85% of gas	14.7 ± 0.4	305 ± 14	5.3	8 steps/8 (1–8) 100% of gas	14.8 ± 0.2
Section 1/frontal part	L90	32°23'45.8"N 79°33'30.0"E	SMA	14.5 ± 0.6	7 steps/9 (1–7) 98% of gas	14.5 ± 0.4	308 ± 40	1.6	6 steps/8 (1–6) 98% of gas	14.7 ± 0.2
Section 1/~10 km from front	L86	32°23'45.8"N 79°33'30.0"E	Plateau	15.4 ± 0.5	5 steps/7 (1–5) 98% of gas	15.4 ± 0.4	325 ± 34	2.7	7 steps/7 100% of gas	15.4 ± 0.2
Section 2/frontal part	C38	32°25'01.1"N 79°42'09.4"E	SMA	12.6 ± 0.2	6 steps/9 (1–6) 98% of gas	12.7 ± 0.2	297 ± 4	1.0	8 steps/8 100% of gas	12.6 ± 0.1
Section 3/frontal part	C30	32°23'28.9"N 79°43'35.0"E	Plateau	12.0 ± 0.4	5 steps/8 (2–6) 93% of gas	12.2 ± 0.2	241 ± 24	0.3	6 steps/8 (1–6) 93% of gas	11.9 ± 0.1
Section 3/~500 m from front	C31	32°23'17.6"N 79°43'25.3"E	Plateau	12.1 ± 0.2	4 steps/7 (2–5) 86% of gas	12.0 ± 0.4	332 ± 44	3.0	5 steps/7 (1–5) 92% of gas	12.4 ± 0.3
Section 4/~3.5 km from front	C43	32°19'34.9"N 79°44'21.1"E	Plateau	12.0 ± 0.4	4 steps/6 (1–4) 97% of gas	12.0 ± 0.4	304 ± 22	0.8	4 steps/6 (1–4) 97% of gas	12.0 ± 0.2

^aErrors are quoted at 2σ confidence level.

of the samples. Most corresponding inverse isochron diagrams yield relatively scattered data (Figures 9 and S1–S10), preventing any reliable isochron age calculations (MSWD > 30). Only for sample C38 an HT isochron age of 17 ± 1 Ma can be calculated ($^{40}\text{Ar}/^{36}\text{Ar}_i = 420 \pm 90$) that is not statistically different from the HT Simple Mean Age (Table 1). The high-temperature steps of the P32, P56 and P57 K-feldspar age spectra have partially scattered ages and do not present reliable plateau or pseudoplateau ages; it was thus impossible to get reliable age estimates or perform thermal modeling on those samples. Locally, older spikes appear on all the increasing middle temperature part of the spectra, but they are relatively minor and were neglected for modeling.

6.2. Micas

[38] Muscovite and biotite age spectra are characteristically flat for most of the gas release (Tables 4 and 5 and Figures 9b and 9c and auxiliary material Figures S11–S17 and S12–S22). Inverse isochron diagrams reveal that the $^{40}\text{Ar}/^{36}\text{Ar}_i$ is systematically close to the atmospheric composition; consequently, inverse isochron ages, plateau ages, and total fusion ages are identical.

[39] Like LT feldspar ages, mica ages are diachronic between the northernmost section 1 and the other sections. In valley 1, white mica ages are 15.4 ± 0.3 to 14.4 ± 0.3 Ma, whereas biotites recorded an age of 14.3 ± 0.3 Ma; micas from other sections (sections 2, 3, 4, 5, and 6) provide constant ages lying in a narrow range of 12.7 ± 0.2 to 10.4 ± 0.8 Ma for muscovite, and 11.9 ± 0.3 to 10.8 ± 0.7 Ma for biotite.

6.3. (U-Th)/He on Apatite

[40] Three replicates, each containing four apatite grains from sample P32 (section 5), were analyzed. Replicates gave relatively good reproducibility (the standard deviation represents less than 14% of the apatite age). Age estimates are comprised between ~4.8 and 3.7 Ma with a mean value of 4.4 Ma (Table 5). Calculated mean error is 0.4 Ma and analytical uncertainties <0.1 Ma (1σ). In the following, only the mean error value based on the samples' reproducibility will be used.

6.4. Fission Tracks on Apatite

[41] Samples P18 and C32, from the frontal part of valleys 2 and 5, yield ages of 5.4 ± 0.4 and 5.5 ± 0.6 Ma, respectively, whereas sample L87, located 10 km south-westward of the frontal part of section 1, shows an age of 8.5 ± 0.5 Ma (Table 6). Pooled ages are reported as all samples pass the chi-squared test. Sample L87 yielded 25 confined track lengths; although normally a larger number are utilized for modeling, an average of 3 tracks Ma^{-1} is sufficient to permit moderate resolution thermal modeling using the AFTSolve model [Ketcham *et al.*, 1999]. The kinetic characteristics of the apatite in this sample were constrained using etch pit diameter measurements (Dpar), calibrated against reported Durango apatite [Ketcham *et al.*,

Table 5. Biotite ⁴⁰Ar/³⁹Ar Results^a

Section/Site	Sample		Plateau Age			Inverse Isochron Age				Total Fusion Age, Ma
	Number	Location	Statistics	Age, Ma	Steps	Age, Ma	⁴⁰ Ar/ ³⁶ Ar _i	MSWD	Steps	
Section 1. frontal part	L91	32°28'46.9"N 79°36'45.0"E	Plateau	14.2 ± 0.3	4 steps/7 (1-4) 98% of gas	14.3 ± 0.2	299 ± 28	1.1	5 steps/6 (1-5) 98% of gas	14.2 ± 0.4
Section 4. ~3.5 km from front	C43	32°19'34.9"N 79°44'21.1"E	Plateau	11.6 ± 0.1	6 steps/9 (1-6) 98% of gas	12 ± 0.2	250 ± 28	0.9	6 steps/9 (1-6) 99% of gas	11.6 ± 0.2
Outcrop 5	P56	32°08'14.7"N 79°56'12.9"E	Plateau	11.8 ± 0.1	7 steps/9 (1-7) 98% of gas	11.8 ± 0.2	311 ± 86	1.3	5 steps/9 (1-5) 98% of gas	11.9 ± 0.2
Outcrop 5	P57	32°08'14.7"N 79°56'12.9"E	Plateau	11.9 ± 0.1	7 steps/10 (1-7) 94% of gas	11.9 ± 0.2	316 ± 60	3.3	7 steps/10 (1-7) 94% of gas	12.2 ± 0.2
Section 6. frontal part	P32	32°01'38.4"N 80°00'38.2"E	Plateau	10.9 ± 0.2	3 steps/10 (2-4) 40% of gas	11.0 ± 0.2	311 ± 8	0.7	4 steps/10 (1-4) 99% of gas	11.1 ± 0.1

^aErrors are quoted at 2σ confidence level.

Table 6. Apatite Fission Track Results^a

Sample	Elevation, m	Location	Spontaneous			Induced			Dosimeter			Pooled					
			Rho-S, × 10 ⁴	No Xls	Rho-I, × 10 ⁶	NS	Rho-I, × 10 ⁶	NI	Rho-D, × 10 ⁶	ND	Pχ ² , %	Age, Ma	±1σ	Length	±2σ	SD	n
P32	4700	32°01'38.4"N 80°00'38.2"E	0.0507	20	2.274	82	1.3272	3680	5291	100	5.5	0.6	21			0	
L87	5230	32°23'45.8"N 79°33'30.0"E	0.1600	23	4.569	350	1.3194	9995	5291	70	8.5	0.5	43	14.3	0.42	1.1	25
P18	4500	32°23'28.9"N 79°43'35.0"E	0.1023	20	4.548	228	1.3038	10137	5291	54	5.4	0.4	44	13.1	2.00	2.7	7

^aThe pooled age is reported for all samples as they pass the chi-square test; error is one sigma, calculated using the zeta calibration method [Hurford and Green, 1983] with zeta of 369.6 ± 7.6 (E. Sobel, unpublished statistical value, 2006). No Xls is the number of individual crystals dated. Rho-S, Rho-I, and Rho-D are the spontaneous, induced, and dosimetry track density measured (tracks cm⁻²). NS and NI are the number of spontaneous and induced tracks counted. Pχ² (%) is the chi-square probability [Galbraith, 1981; Green, 1981]. Sd and n are the standard deviation of the track lengths and number of measurements, respectively. Values greater than 5% are considered to pass this test and represent a single population of ages. ND is the number of tracks counted in determining Rho-D; CN5 dosimetry glass used.

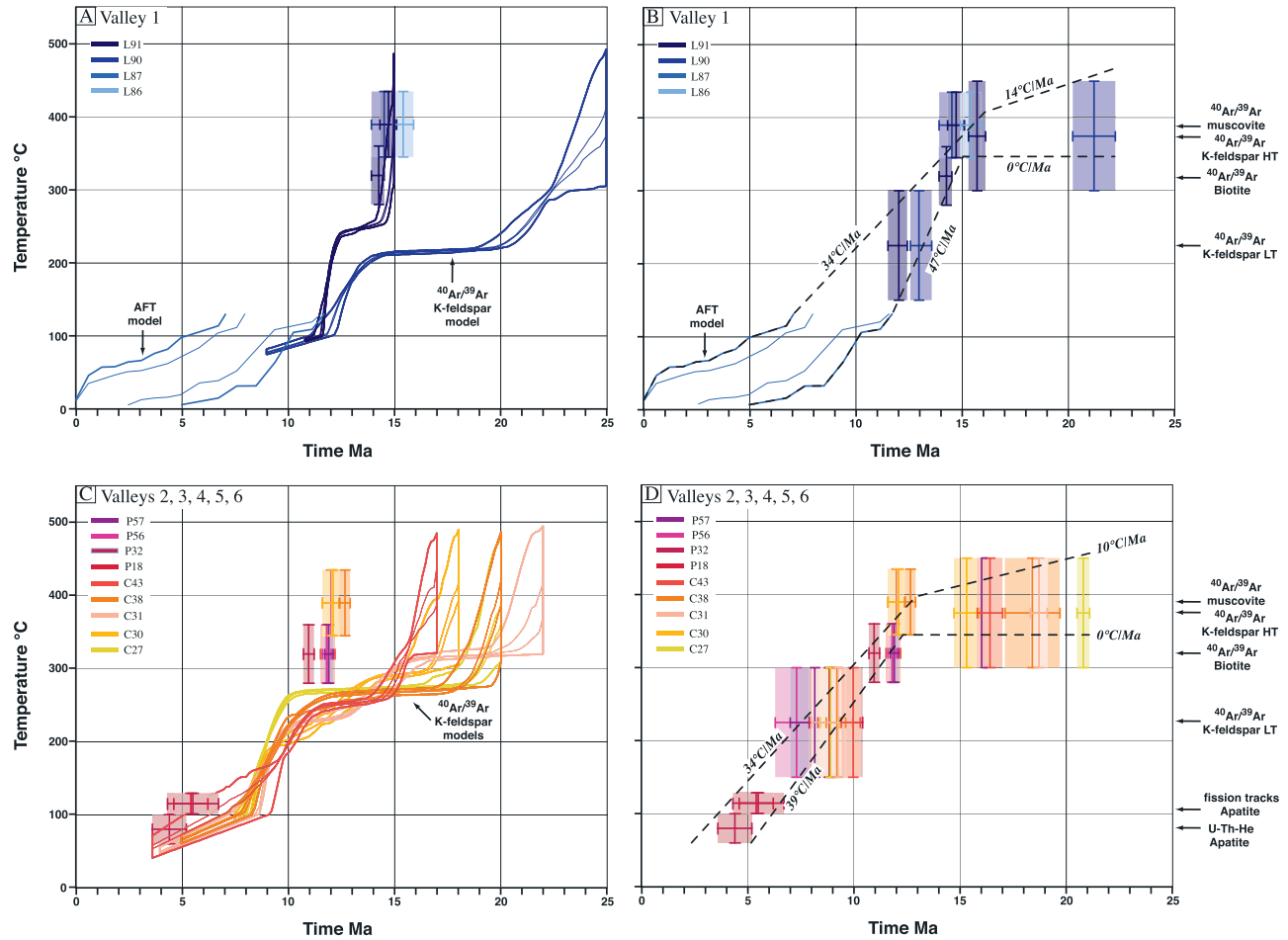


Figure 10. Time-temperature diagrams showing biotite, muscovite, and K-feldspar $^{40}\text{Ar}/^{39}\text{Ar}$ age estimates, together with fission track (FT) and (U-Th)/He results on apatite. Ages are systematically $\sim 2\text{--}3$ Ma older in the northernmost valley (section 1) than in the other sections (sections 2, 3, 4, and 5). This implies a different time-temperature evolution between the northernmost section and the others, and consequently, data from this valley are plotted separately. (a and b) Northern most field section (section 1, location on Figure 2). (c and d) Other sections (sections 2, 3, 4, 5, and 6). Age uncertainties for the $^{40}\text{Ar}/^{39}\text{Ar}$ analysis include the entire range of isochron and plateau age estimates each quoted at the 2σ confidence level; fission track and (U-Th)/He age estimates at 2σ confidence level (see text for details). Continuous lines represent K-feldspar (Figures 10a and 10c) and apatite cooling models (Figures 10a and 10b, see text for details). Thick and fine curves are the 95 and 65% confidence intervals, respectively, of the best fit cooling history (K-feldspar modeling has tested various solutions by a Monte Carlo algorithm to assess the variance of the resulting best fits). In Figures 10a and 10c, note the inconsistency between most K-feldspar thermal models and mica age estimates. In Figures 10b and 10d, preferred time-temperature history for the Ayilari range (see text for explanations). The cooling history is most likely bounded by the black dashed lines, which were drawn by using a minimum of straight segments passing through all the data at 2σ confidence level.

1999]; a value of $1.55\ \mu\text{m}$ was used. The T_A for this sample is $\sim 130^\circ\text{C}$.

7. Constructing a Cooling History

7.1. Closure Temperatures and Cooling Paths

[42] Once the age of a given mineral is calculated, a fundamental and controversial issue is to determine whether this age corresponds to mineral crystallization, recrystallization,

or cooling below a given closure temperature. The temperatures estimates for ductile deformation are above $350\text{--}400^\circ\text{C}$; higher than closure temperatures for the radiogenic systems used in this study. Therefore $^{40}\text{Ar}/^{39}\text{Ar}$, AFT and (U-Th)/He ages most likely correspond to cooling ages, and we choose to systematically build cooling histories from our results. This choice leaves open further discussion on the internal consistency of such cooling histories and the relationship between $^{40}\text{Ar}/^{39}\text{Ar}$ ages and deformation

events. For example, $^{40}\text{Ar}/^{39}\text{Ar}$ ages of the finest white micas may in fact correspond to the age of synkinematic crystallization [e.g., *Kirschner et al.*, 1996]. We assumed closure temperature of $390 \pm 45^\circ\text{C}$ for the white micas [*Hames and Bowring*, 1994, and references therein], $320 \pm 40^\circ\text{C}$ for biotites [e.g., *Harrison et al.*, 1985], $115 \pm 15^\circ\text{C}$ for the AFT (when thermal modeling was not possible [*Gleadow et al.*, 1986; *Ketcham et al.*, 1999]), and $80 \pm 20^\circ\text{C}$ for the (U-Th)/He on apatite [*Farley*, 2000].

[43] Most K-feldspar data permitted modeling of their cooling history. We restricted our models to the simplest possible solution: monotonic cooling. Allowing transit reheating would lead to significantly different but less constrained thermal histories. Modeling was performed on 7 K-feldspar fractions. Modeled age spectra are shown on Figures 9 and S1–S10 and cooling histories are presented on Figures 10a and 10c. The experimental Arrhenius diagrams were satisfactorily fit with diffusion models assuming 5 to 8 domains sharing common activation energy on the order of $33.5\text{--}39.9 \text{ kcal mol}^{-1}$. The best cooling models invariably show two rapid cooling phases separated by an isothermal or slow cooling rate period. The age of the oldest cooling event varies over a wide range of $\sim 23\text{--}14 \text{ Ma}$ in relation with the variability of HT ages in K-feldspars, but these ages are poorly constrained, since it was impossible to obtain a perfect fit for the oldest part of the age spectrums (Figures 9 and S1–S10). Moreover the modeling is based on the assumption that argon diffuses out of the lattice, while within the furnace, breakdown of the feldspars takes place in the range $1050\text{--}1100^\circ\text{C}$, thus violating the model hypothesis above these temperatures. Modeling above 1100°C is therefore unconstrained. Following cooling from high temperature a period of slow cooling rate ($<15^\circ\text{C Ma}^{-1}$) is modeled and corresponds to a drop of temperature down to $\sim 210\text{--}260^\circ\text{C}$. The youngest cooling event is always very well constrained at $\sim 12 \text{ Ma}$ for samples of the northernmost valley (section 1) and at $\sim 11\text{--}9 \text{ Ma}$ for all the others (sections 2, 3, and 4).

7.2. Pertinence of the Cooling Models From K-Feldspars

[44] The time-temperature diagram reveals an incompatibility between the K-feldspar time-temperature modeling and the rest of the data set: all but L91 K-feldspar cooling curves, and especially the protracted slow cooling periods, do not match the mica ages if classical mica closure temperature are accepted (Figures 10a and 10c). It is noteworthy that the white mica ages of dextrally deformed samples (L90, L91) from valley 1 are identical to those of white micas coming from a nondeformed sample (L86) located in the same valley. Similarly, consistent biotite ages determined from deformed samples of valleys 4 and 6 are contemporaneous with nondeformed biotite ages from samples located in valley 5 (P57, P56). This strongly suggests that the micas crystallized prior to the last deformation increments, and thus not below their closure temperature, ruling out exceptionally low temperatures that would bring the mica ages down on the feldspar cooling curve. Therefore micas $^{40}\text{Ar}/^{39}\text{Ar}$ ages most likely represent cooling

ages at “normal” closure temperature. Similarly, there is no reason to suspect abnormally high kinetic characteristics in the feldspars that would imply that the less retentive domains closed at higher temperature around $11\text{--}9 \text{ Ma}$. Our calculated activation energies $33.5\text{--}39.9 \text{ kcal mol}^{-1}$ are on the low side compared with the statistical study of *Lovera et al.* [1997], which showed that feldspars from various origins and terrains have activation energies varying around a mean of $\sim 46 \pm 6 \text{ kcal mol}^{-1}$. However, even activation energies of 50 kcal mol^{-1} would not imply closure temperatures high enough to reconcile feldspars and mica data.

[45] On the other hand, petrostructural observations show significant synkinematic recrystallization of K-feldspar with the development of a mantle of recrystallized grains around crystal cores. Thus the dated K-feldspars were most likely constituted of clusters made of several generations of intimately mixed crystals having grown, crystallized and cooled at various times during the deformation rather than monogenetic single crystals with simple diffusion domains. This strongly suggests that thermal models are not reliable in that case.

[46] Since thermal modeling using simple diffusion and monotonic cooling does not adequately explain the continuous recrystallization of the dated phases, we therefore used a cruder approach, assuming accepted closure temperatures for micas. Since pseudoplateaus exist for both LT and HT ages, it is likely that argon loss was limited at the corresponding times. This implies that rapid cooling took place at those periods, pointing to a limited closure temperature window at those times. We also make the hypothesis that, although the kinetic characteristics calculated from the continuous thermal modeling of feldspars are not adequate to be used for a single grain, they can be used as crude estimates for the closure temperatures of crystals degassing during LT and HT phases. Using these data suggest that the LT ages correspond to cooling at $\sim 150\text{--}300^\circ\text{C}$, whereas HT ages (at $1000\text{--}1400^\circ\text{C}$ furnace temperatures) correspond to cooling at $\sim 300\text{--}450^\circ\text{C}$ (Figures 10b and 10d) [e.g., *Leloup et al.*, 2001]. As it is often difficult to assess which of the isochron or plateau ages is the most realistic because both are equally acceptable statistically, we quote uncertainties that include the entire range of isochron and plateau age estimates, the resulting global uncertainties are taken at the 2σ confidence level.

8. Geochronological Constraints on the Age of Right-Lateral Shearing and Normal Faulting Along the Ayilari Range

8.1. High- to Medium-Temperature Thermochronological Data (Argon HT Kf) and (U-Th/Pb) Geochronology: Evidence for a First Lower Miocene Rapid Cooling Event

[47] U-Th-Pb ages obtained by classical Isotopic Dilution–Thermoionisation Mass Spectrometry (ID-TIMS) and by Secondary Ion Mass Spectrometry (SIMS) on zircons and monazites [*Lacassin et al.*, 2004; *Valli*, 2005, also manu-

script in preparation, 2007] reveal that the most recent magmatic episode took place at ~25–21 Ma in the Ayilari range. This episode was synkinematic with right-lateral shear and was accompanied and followed by metamorphism from ~22 to 19 and hydrothermalism between ~22 and 15 Ma [Valli, 2005, also manuscript in preparation, 2007] (Figure 11a). These results are interpreted as reflecting heat and fluid advection during deformation along the KFZ. Those ages are very close to or identical within error to

the HT ages from K-feldspars of the present study. Theoretically, the granitoid crystallization temperature (~750–800°C [Clemens, 2003]) is ~350°C above the feldspar HT cooling temperature (~400°C) (Figure 11), which should have lead to significantly different ages. Three hypotheses can be envisaged to explain both data sets.

[48] 1. K-feldspar HT ages could be too old because of excess argon. However, a close analysis of inverse isochrons show that there is no clear evidence for such excess;

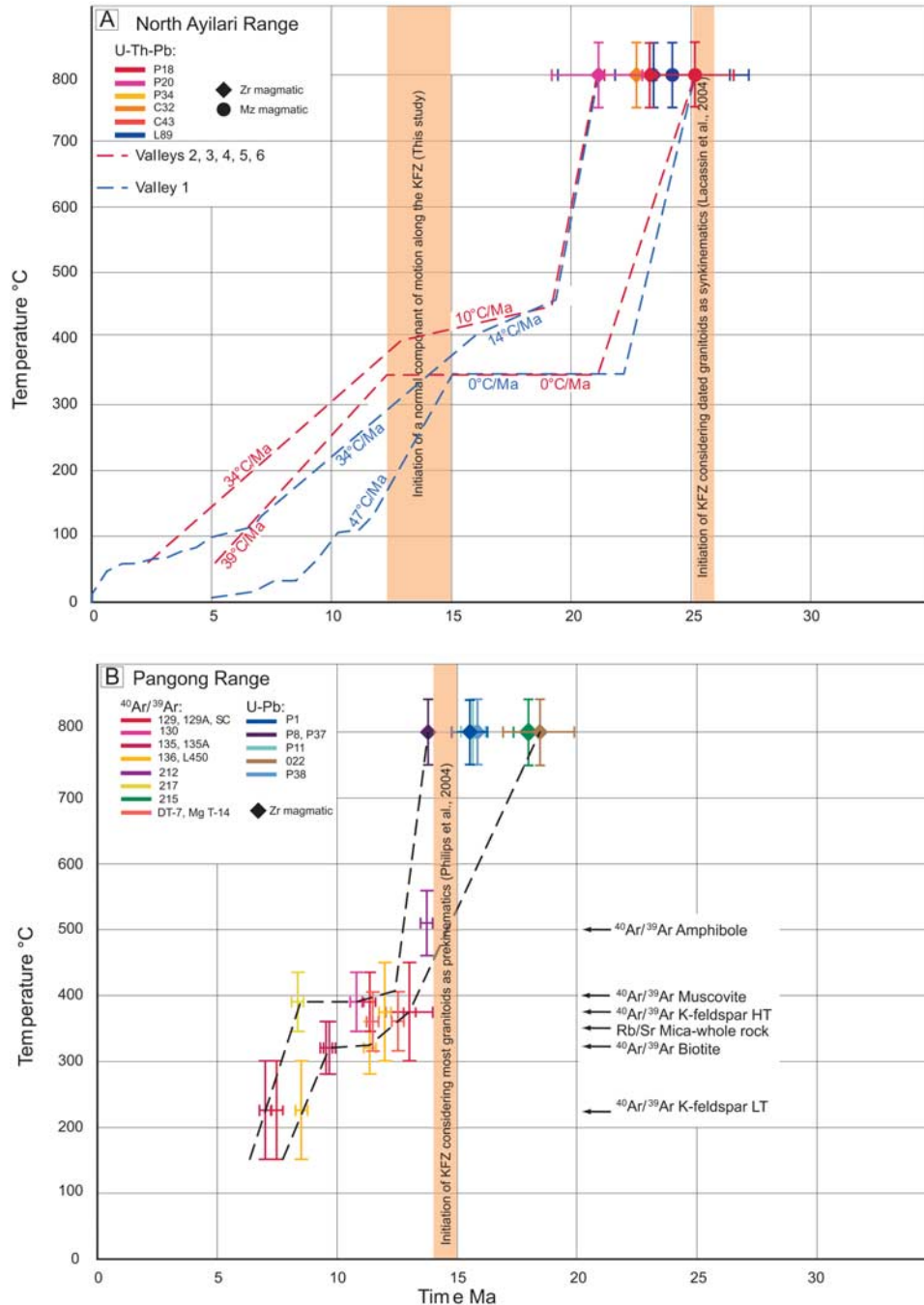


Figure 11

the age of the only sample with an HT isochron age is confirmed by the simple mean age (see section 6.1). This strongly suggest that little excess argon is present and that K-feldspar HT ages are meaningful as is the spread in age between samples of ~ 6 Ma observed in each section (Figure 10).

[49] 2. K-feldspar HT ages could correspond to unusual high closure temperatures. However, as discussed in section 7, there is no reason to suspect such a case.

[50] 3. Both magmatic zircon U/Pb and argon HT K-feldspars ages are correct and associated with the expected closure temperatures. This implies that the early Miocene (~ 25 – 21 Ma) synkinematic magmatic episode was immediately followed by a period of rapid cooling during which metamorphism took place until ~ 19 Ma and metasomatism until ~ 15 Ma (Figure 11). For example, in section 1, temperature dropped from solidus temperature ($< 750^\circ\text{C}$ for two-mica granitoids [Gardien *et al.*, 1995]) at 24.2 ± 2.4 Ma (L89 magmatic monazite), to closure temperatures corresponding to the HT part of the K-feldspars age spectra ($400 \pm 50^\circ\text{C}$) at 18.4 ± 1.3 (C38, most reliable KfHT age) (Figure 11a). Because hydrothermal monazites and zircons can have crystallized in a large range of temperatures (~ 350 – 700°C [Seydoux-Guillaume *et al.*, 2002]), it is not possible to draw more precise cooling histories. For example, L89 metamorphic or hydrothermal monazites have the same age than the HT K-feldspar.

[51] Rapid cooling soon after 25–21 Ma could result from tectonic exhumation immediately after the granitoids emplacement and associated with thrust or normal components of displacements along the KFZ at that time. However, in the North Ayilari range all kinematic indicators associated with deformation under high- to medium-temperature conditions (> 600 – 400°C) indicate an almost pure dextral shear with horizontal lineations (see section 3.1. and Figure 5). Dextral-normal deformation is restricted to the more brittle and recent stages of deformation (see below). A phase of tectonic exhumation coeval with this early Miocene rapid cooling event is thus unlikely. Alternatively, rapid cooling could follow high temperature shearing at shallow depths. High geothermal gradient due to shear heating and/or heat advection is common in strike-slip shear zones [e.g., *Leloup et al.*, 1999] It may induce fast local

cooling when deformation slows down or ceases at one location, even if it continues in another part of the shear zone. Such process could explain the differences between the various HT K-feldspar ages, as the samples are located at different places within the shear zone. Unfortunately, we do not have good constraints on the depth of right-lateral deformation in the North Ayilari range to confirm this hypothesis.

[52] In the Pangong range 200–350 km to the northwest (Figure 1) outcrop other right-laterally sheared ductile mylonites [Searle *et al.*, 1998; Dunlap *et al.*, 1998; Rolland and Pêcher, 2001; Rolland, 2002; Rolland *et al.*, 2007]. Dunlap *et al.* [1998] state that the strike-slip motion is restricted to amphibolite and greenschist metamorphic conditions. However, recent petrologic and thermobarometric studies constrain the P-T-t path during this deformation [Rolland and Pêcher, 2001; Rolland *et al.*, 2007]. Both granulitic ($\geq 800^\circ\text{C}$ and ~ 5.5 – 5 kbar) and amphibolitic (650°C and 6–7 kbar) conditions occur. Peak temperature in the amphibolitic rocks were attained at ~ 18 Ma, at the time of Tangtse granite emplacement (Figure 11b, sample 450 of Rolland *et al.* [2007]). As strike-slip deformation continued, temperature dropped to $\sim 500^\circ\text{C}$ at ~ 13.6 Ma (Figure 11b, sample 441 of Rolland *et al.* [2007]) during cooling to low T amphibolitic conditions (500 – 300°C and 2–5 kbar). This suggests that strike-slip deformation at middle to upper crustal levels (depth ≤ 20 km) occurs under anomalously high temperature conditions. The shear zone thermal state was probably very similar in the North Ayilari range.

8.2. Medium- to Low-Temperature Cooling History of the Ayilari Range: Evidence for a Second Rapid Cooling in the Middle Miocene

[53] Below 400°C , time-temperature paths can be drawn from argon (muscovite, biotite, LT K-feldspar), apatite fission track and apatite (U-Th)/He data (Figures 10b, 10d, and 11a). These paths show good agreement between the different thermochronometers since, within error, age estimates decrease with the closure temperature of the minerals (Figures 10b and 10d).

[54] At temperature around 400°C , the HT parts of the K-feldspars spread in age between ~ 22 – 15 Ma in section 1

Figure 11. Cooling histories for the Ayilari and Pangong Ranges. Dashed lines represent the cooling histories. Mag, magmatic; Zr, zircon; Mz, monazite. Ages are plotted with a 2σ confidence level. (a) North Ayilari range. U-Th-Pb data constraining the high temperature part of the diagram are from the work of F. Valli *et al.* (manuscript in preparation, 2007), whereas the medium- to low-temperature evolution is documented in this study (Figure 10). The zircons and monazites are taken as constraining the crystallization of the syntectonic granites at ~ 750 – 800°C [Clemens, 2003] The timing of onset of the KKF in the Ayilari range [Lacassin *et al.*, 2004] and the onset of normal faulting (this study) are shown. (b) Pangong Range. Data compiled from the literature and discussed in the text. Samples 129, 129A, 130, 135, 135A, 136, 212, 215, and 217 are from the work of [Dunlap *et al.*, 1998]; the samples P1, P8, P11, P37 and P38 are from [Phillips *et al.*, 2004]; and the samples 450, 022, and SC are from Searle *et al.* [1998]. Data from U/Pb analysis have been plotted accordingly to the published ages with uncertainties as discussed in the original papers, the temperature of crystallization being arbitrarily chosen between 650°C and 800°C as postulated by F. Valli *et al.* (manuscript in preparation, 2007). Data from K-feldspar analysis were treated in a similar way to those from the present study: from the original data low-temperature (LT) and high-temperature (HT) pseudoplateaus were appreciated and are plotted with a range of closure temperature of 150 – 300°C and 300 – 400°C , respectively (see text for discussion). Ages are plotted with a 2σ confidence level.

(Figure 10b) and ~ 21 – 15 Ma in the southern valleys (Figure 10d), while muscovite ages are grouped between ~ 15 – 14 Ma and ~ 13 – 12 Ma, respectively (Figures 10b and 10d). These suggest cooling rates lower than 14°C Ma^{-1} during the ~ 22 – 14 Ma period in section 1, and lower than 10°C Ma^{-1} during the ~ 21 – 12 Ma period in the southern valleys. The flatness of mica plateaus (Figures 9b, 9c, and S11–S22) is atypical for gneisses or granitoids; it is more characteristic of rocks in which crystals were quenched. This, together with the narrow age difference between muscovite and biotite, strongly suggests the initiation of rapid cooling at ~ 15 – 14 Ma in the northernmost section, and at ~ 13 – 12 Ma in the other sections. The thermochronometric estimates, in the time period ~ 15 – 5 Ma for section 1 and ~ 13 – 4 Ma for the other sections, suggest monotonic rapid cooling at an average rate of $\sim 40^\circ\text{C Ma}^{-1}$ and $\sim 36^\circ\text{C Ma}^{-1}$, respectively (Figures 10b and 10d). This second episode of rapid cooling is clearly associated with petrostructural indicators of medium to low temperature, indicating the onset of a component of normal throw on the Karakorum fault in the Gar basin. This kinematic change induced exhumation and faster cooling and was probably associated with the rise of the Ayilari range.

[55] In summary petrostructural data not only show that deformation with a dextral component of shear started at temperature >400 – 600°C but also that it has been continuous up to near-surface conditions, and was thus associated to the cooling history deduced in the 21 to ≤ 4 Ma age range. Our thermochronological results reinforce and extend our previous conclusions [Lacassin et al., 2004; Valli, 2005; F. Valli et al., manuscript in preparation, 2007], now demonstrating not only that right-lateral motion was already in progress in the early Miocene (25–21 Ma) but also that right-lateral shear with a normal component initiated at 12–15 Ma and continued until at least 4 Ma. Since such deformation is still active [e.g., Armijo et al., 1989; Chevalier et al., 2005], this implies at least 20 million years of deformation along the southern part of the Karakorum Fault zone.

9. Large-Scale Geometry and Kinematics of the Karakorum Fault Zone

9.1. Large-Scale Kinematics of the KFZ

[56] Along the entire KFZ, the crystallization ages of numerous granitoids which are likely synkinematic, as well as thermochronological ages ($^{40}\text{Ar}/^{39}\text{Ar}$, fission tracks, (U-Th)/He) of cooling coeval with dextral shear, also suggest an early onset of deformation along the KFZ.

[57] As already mentioned, dextrally sheared mylonites of the KFZ also outcrop in the Pangong range and in the Nubra valley (Figure 1). In the Pangong range, a ~ 6 km wide band of mylonitized leucogranites, migmatites, amphibolitic and granulitic gneisses outcrop in between two strands of the KFZ [Searle et al., 1998; Rolland et al., 2007]. U/Pb shrimp zircon ages of 18.0 ± 0.6 (sample 215) and 18.5 ± 1.5 (sample O22) were interpreted as dating crystallization of the Tangtze granite [Searle et al., 1998].

Mylonitic leucogranites from Tangtze and Nubra yield U/Pb ages of ~ 16 Ma (samples P1, P11, and P38 [Phillips et al., 2004]), while crosscutting dikes yield ages of ~ 14 Ma (samples P8 and P37 [Phillips et al., 2004]) (Figure 11b). Such ages are interpreted to bracket the time of initiation of right-lateral faulting along the KFZ because ductile deformation is considered to have started after granite crystallization and have ended prior to leucocratic veins emplacement [Phillips et al., 2004]. While such ages dismiss the idea of a purely Holocene KFZ, as proposed by Searle [1996] on a regional basis, they imply a younger onset than the ≥ 21 Ma age proposed in the Ayilari range [this study; Lacassin et al., 2004; F. Valli et al., manuscript in preparation, 2007].

[58] Several lines of evidence suggest that the Tangtze granite is synkinematic to the KFZ and thus that ~ 18 and ~ 16 Ma are only lower bounds, not upper bounds, for the age of this fault.

[59] 1. As already emphasized for the North Ayilari range (section 2-2, Figure 5d) the emplacement of a network of variously deformed dikes which come together to form a main sheared pluton represent reliable field evidence of synkinematic crystallization [Lacassin et al., 2004]. Such field relationships correspond to the Tangtze granite and associated dikes [Phillips et al., 2004; Searle et al., 1998; Weinberg and Searle, 1998].

[60] 2. S-C fabrics are ubiquitous in the North Ayilari range (Figures 5b and 5c), and in the Tangtze granite [Rolland, 2000; Searle et al., 1998]. Such structures are typical of shear deformation under high to medium-grade temperature conditions, and they might be typical of synkinematic granitoid [e.g., Gapais, 1989a, 1989b; Gapais and Barbarin, 1986]. The deformed rocks indeed show clear evidence for the re-crystallization of feldspar and quartz, the latter by subgrain rotation and grain boundary migration [Searle et al., 1998, Figure 8; Rolland, 2000], demonstrating that strike-slip deformation occurred at temperature above 600 – 400°C (see section 2–3). Such temperatures are close to the 750°C solidus temperature of two mica-bearing granites and dikes [Gardien et al., 1995, and references therein]. This suggests that crystallization of the Tangtze granite might have been synkinematic. The possibility of high-temperature (>350 to $\leq 750^\circ\text{C}$) strike-slip deformation was already evoked by Dunlap et al. [1998, p. 904], and Weinberg and Searle [1998, pp. 885 and 890].

[61] 3. The Tangtze granite is intrusive within granulitic and amphibolitic rocks. The P-T path of these rocks is retrograde from granulitic ($T > 800^\circ\text{C}$, $P \sim 5.5$ kbar) or amphibolitic ($T \sim 750$ – 700°C , $P \sim 4$ – 5 kbar) to greenschist ($T \sim 500^\circ\text{C}$, $P \sim 3$ kbar) conditions [Rolland and Pêcher, 2001; Rolland et al., 2007]. These host rocks were penetratively deformed in a dextral transpressive environment at temperature ranging from $>800^\circ\text{C}$, thus above the Tangtze granite solidus temperature (750°C), to surface conditions. This again suggests a synkinematic emplacement of the granite.

[62] One may speculate that other granitoids along the fault may also be used as evidence for syntectonic magmatism and metamorphism. Farther to the NW, the Baltoro batholith crystallized between 26 and 20 Ma [Parrish and

Tirrul, 1989; Schärer *et al.*, 1990], at relatively low temperature (750–600°C) [Searle *et al.*, 1992]. The batholith has a sigmoidal shape; the southern and northern edges strike \sim N110°, 100 km west of the KFZ and progressively bends eastward to become parallel to the KFZ (N142° strike, Figure 1). South directed thrusting occurred during granite emplacement along the southern edge of the batholith [Searle *et al.*, 1992]. The eastern border of the batholith is crosscut by the KFZ active strand and shows dextral S-C structures [Searle *et al.*, 1998]. Similar to the Tangtse granite, the Baltoro batholith has been interpreted as predating the onset of dextral motion along the KFZ [Searle, 1996; Searle *et al.*, 1998]; however, its large-scale sigmoid shape suggests an intrusion synchronous with right-lateral deformation along the KFZ, coeval with thrusting on the batholith southern edge and dextral high to medium-temperature deformation along its eastern border. In this case, the 26–20 Ma ages of the batholith would, as already suggested by Mahéo *et al.* [2004], represent a lower bound for the initiation of the KFZ.

[63] In the Pangong range, Dunlap *et al.* [1998] report two phases of rapid cooling, one at circa 17–13 Ma followed by an intervening period of relatively slow cooling, and the second starting at circa 8 Ma and continuing until at least 7 Ma [Dunlap *et al.*, 1998] (Figure 11b). This time-temperature evolution was mainly based on K-feldspar models, which, as in the Ayilari range, are barely consistent with micas ^{40}Ar – ^{39}Ar ages from the same samples, or from the same structural unit. This suggests that these K-feldspar thermal models are no more reliable than ours. However, a cautious interpretation of the data, assuming that the mica and amphibole ^{40}Ar – ^{39}Ar ages correspond to cooling ages, shows that a single monotonic cooling history cannot perfectly reconcile the entire data set (Figure 11b), as is the case in the Ayilari range, suggesting that the shift in cooling rates at circa 8 Ma is real. This rapid cooling event occurred during right-lateral transpression [Dunlap *et al.*, 1998; Rolland, 2000, 2002; Rolland and Pêcher, 2001], whereas the contemporaneous rapid cooling event recorded in the Gar-Baer basin (between \sim 16–12 and $<$ 4 Ma) is associated with transtension [this study; Lacassin *et al.*, 2004; F. Valli *et al.*, manuscript in preparation, 2007].

[64] The shift from transpression to transtension along the fault could be explained by “zipper tectonics” [Harrison *et al.*, 1996], a model thoroughly documented along another major continental strike-slip fault: the Ailao Shan–Red river shear zone [Leloup *et al.*, 2001]. If the KF was a transform fault in the strictest sense, it would follow a small circle of the Euler rotation pole between the Ladakh and Qiangtang blocks. Motion would be purely strike slip, and there would be only local denudation along the fault zone in restraining or releasing bends. On the other hand, if the fault does not follow a small circle, transpression and transtension would take place along large portions of the faults; in the NW and SE parts of the right-lateral KFZ if the rotation pole is located NE of the fault. The neutral point that corresponds to the place where the fault zone aligns with a small circle is the only place where motion is purely strike slip and would be located somewhere between the

Pangong and Ayilari ranges, thus allowing transtension in the Gar-Baer basins and transpression in the Pangong range and to the north. Since these ranges underwent a strike-slip deformation with only minor denudation prior to 18–12 Ma, a regional kinematic change may have occurred at that time, possibly related to blocks readjustments linked to the initiation of the left lateral Gozha fault which formed the northern boundary of the Pangong range (Figure 1). It should be emphasized that this model remains highly hypothetical because of the scarcity of well constrain time-temperature paths along the KF and within adjacent blocks.

9.2. Formation of the Gar Pull-Apart Basin and Hints of Long-Term Slip Rate of the KFZ

[65] The exhumation of the Ayilari range by normal faulting along its northeastern edge was probably associated with the development and deepening of the Gar and Baer pull-apart basins concurrent with the rise of the regional topography. This drove the incision and entrenchment of major river courses that flow in these basins, most notably that of the Indus River (Figure 12a). Likely, the gorges that permit to the Indus River to flow across the Ayilari range date back to the time of rapid uplift of this range at 15–12 Ma, as suggested by our thermochronological data. Matching this outflow channel of the Indus with an upstream inflow channel will thus give an offset accrued in the last 12 to 15 Ma. If this offset correspond to the present-day 120 km offset of the Indus River course along the northern branch of the KFZ (Figure 12b) this would imply dextral motion at a minimum long-term average rate of $9 \pm 1 \text{ mm yr}^{-1}$ (Figure 12b). Such a reconstruction also implies that the river now flowing in the Bangong lake was the upstream part of the Shyok river \sim 14 Ma ago.

[66] However, as with most major rivers crossing active faults, the present-day offset probably represents only a lower bound of finite motion along the northern branch of the KFZ [Gaudemer *et al.*, 1989], because the watershed of the Indus River could have been highly modified during the last millions of years. In the Middle Miocene, this river possibly flowed on the southern side of the Kailas range, where the Gangrinboche molasses were deposited [Davis *et al.*, 2004; Gansser, 1964; Miller *et al.*, 2000], implying a much larger offset of \sim 280 km (Figure 12c). In this scenario, after the Indus river started to entrench across the rising Ayilari range (15–12 Ma), the dextral offset of its upstream and downstream channels forced the river to flow along the fault, in a way similar to the actual Gar River. Thrusting in the Kailas area [Yin *et al.*, 1999]. might have started to change the watershed of the Indus River at circa 13 Ma. The actual upstream channel and watershed of the Indus river, which presently flows from the Shiquanhe basin, might have been captured only recently \sim 10 km south of Zhaxigang. Taking into account such a hypothetical scenario (Figure 12c) would bring the estimate of the long-term average rate on the northern branch of the KFZ to $20 \pm 3 \text{ mm yr}^{-1}$.

[67] For comparison, the total long-term offset on the two branches of the KFZ is estimated at 280 to 400 km from the

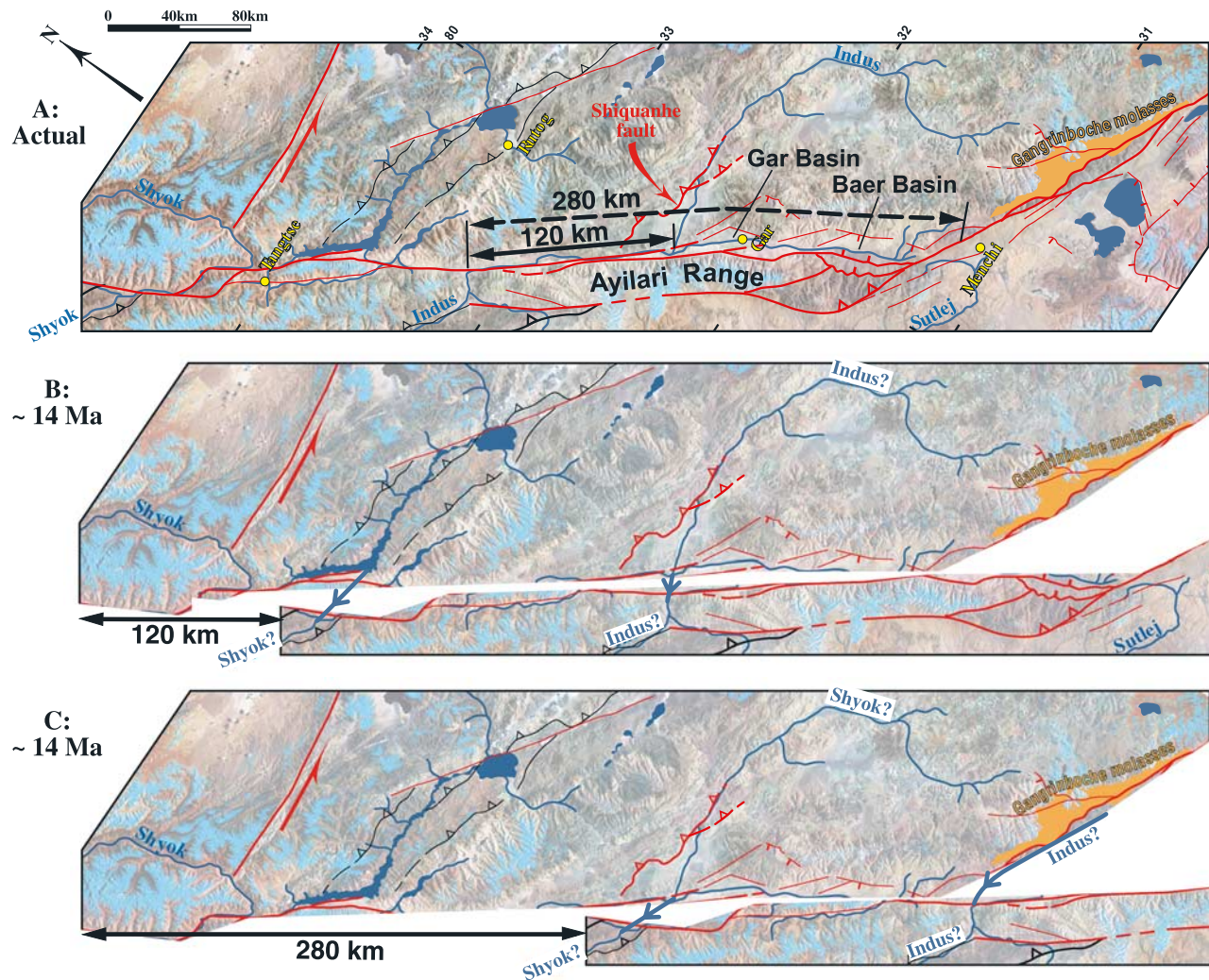


Figure 12. Progressive offset along the Karakorum fault as can be postulated from the entrenchment of the Indus and Shyok rivers in the Ayilari range at ~ 14 Ma (see text for explanations). (a) Arrows represent possible right-lateral apparent offset of the Indus river across the Karakorum fault. (b and c) Corresponding hypothetical reconstructions at ~ 14 Ma if the Shiquanhe basin is taken as the upstream of the Indus or Shyok river at that time. In the latter case the upstream of the Indus had to be somewhere south near Mount Kailas and could be at the origin of the deposits of the Kailas molasses. The Landsat 7 image is projected in UTM 44 using the ellipsoid WGS84.

offset of the Shyok and Indus suture zones [Lacassin *et al.*, 2004], corresponding to average rates of 12 ± 1 to 17.5 ± 1.5 mm yr $^{-1}$ assuming fault initiation between 25 and 21 Ma ago.

[68] **Acknowledgments.** Fieldwork was supported by the programs Intérieur de la Terre (IT) and Dynamique et Evolution de la Terre Interne (DYETI) of CNRS-INSU and by the Institute of Geology, Chinese academy of Geological Sciences, Beijing, China. We thank Mike Murphy, Yann Rolland, and an unknown reviewer for their useful comments, which significantly improved this manuscript.

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